Normal and Differential SAR Interferometry

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Abstract

SAR Interferometry explores the height structure inside a Radar pixel. Two antennas separated with a so called baseline distance B in cross velocity direction take two independent images from the same scene and compare the two phases which result from the same pixel at the two antennas. This phase differences compare to different distances between the Pixel and the two antennas and from this difference the height of the respective pixel in relation to the altitude of the radar over a defined level can be determined.

The Interferogram formation will be explained and the respective terminology will be presented. In this context basic relations and important geometrical and electrodynamics system parameters will be presented and defined, the phase unwrapping will be considered as well as basic interferometry processing procedures and image formation principles. Coherence is the basis for Interferometry as well as a careful calibration. Therefore, the role of interferometric coherence for image interpretation will be shown. Errors caused by phase noise, de-correlation, co-registration misalignment, atmospheric disturbances, baseline limits, etc will be considered. They, principally, are the main reasons for coherence degradation. Different Interferometry modes Cross Track and Along Track Interferometry, single and dual pass Interferometry, Differential Interferometry, and Permanent Scattering Interfererometry SAR will be discussed.

1. Introduction

Goal of this lecture is to present the basics of SAR Interferometry in order to prepare the bottom for the following lectures on Polarimetric Interferometry and its applications. An Electromagnetic Field is characterized by Amplitude E_0, Frequency f, Phase φ and Polarization p, the latter describes the Vector character of the field $\mathbf{E}=(E_X;E_Y)$ principally as functions of time t. E_X and E_Y are orthogonal to each other and characterize the polarization base of a Transversal Electromagnetic wave (TEM wave) which has no component in propagation direction. It is sufficient for a principal understanding to consider in a first step "scalar" Electromagnetic Waves and there behaviour. Tat means without limitation of generality one component of the Electric Field Vector only will be used, and polarization effects, principally, will be excluded. The scalar Electromagnetic Wave is written as:

$$E= E_0 \exp i(2\pi f + \varphi) \qquad (1)$$

2. SAR Interferometry Principles

SAR Interferometry is an established technique for rapid and accurate collection of topographic data, which is essential for establishing Digital Elevation Models (DEM's), it allows the location of the effective scattering centre inside the resolution cell.

Normally, SAR interferometry works with two antennas separated in cross velocity direction characterised with the baseline distance B taking two independent images from the scene observed and comparing the two phases which result from the same pixel. This offers the capability to generate topographic maps. Adding a third flight at the same passes as before and subtracting the two results from each other is called Differential Interferometry.

Due to the distance between the antennas the distance to a single point scatterer on the ground is slightly different by the amount of *r*. This slant range difference *r* corresponds to a phase difference Φ comparing the

received signals of both antennas. After detection of Φ the local terrain elevation above a well known reference area, a geoid for instance, can be determined according to the following equations, which are based on the geometry for SAR interferometry depicted in Fig.1.

 In practice two different methods are applied to obtain an interferometric baseline:
• One pass interferometry: In this case both antennas are mounted rigidly to the same platform with a fixed baseline. There is one transmitter for the two receiving channels. An interferometric image is obtained after one pass of the platform. For the Phase difference results from the two phases φ_1 and φ_2:

$$\varphi_1 = 4\pi \frac{R_1}{\lambda} \ , \ \varphi_2 = 2\pi \frac{2R_1 + r}{\lambda}$$

$$\Phi_{1pass} = \varphi_2 - \varphi_1 = 2\pi \frac{r}{\lambda}$$

• Two pass interferometry: Only one Transmit-Receive channel with one antenna is used, but the platform has to pass the same terrain twice in the same direction along a trajectory parallel to the first one. In this case results for the Phases:

$$\varphi_1 = 4\pi \frac{R}{\lambda} \ , \ \varphi_2 = 4\pi \frac{R+r}{\lambda}$$

$$\Phi_{2Pass} = \varphi_2 - \varphi_1 = 4\pi \frac{r}{\lambda}$$

For the Phase difference Φ will be written in the following:

$$\Phi = m \ 2\pi \frac{r}{\lambda} \tag{2}$$

With m=1 for the one pass case and m=2 for the two pass case. The propagation path difference $2r$ is usually much grater than the wave length, therefore, Φ normally shows an ambiguity of many cycles. In any case it can be determined mod 2π only, the phase is wrapped.

For topography purposes an across-track arrangement of the two sensors is essential. However, along track arrangements, principally, allow the measurements of the ground velocity if targets ore pixels respectively. Both modes are depicted in Fig. 2. The coherence between the two signals is essential in any case.

2.1 Basic Interferogram Relations
From the geometry depicted in Fig. 1 the following basic Interferometry relations result:
$$z = H - R\cos\vartheta \tag{3}$$
After some manipulations results for the slant range:
$$R = \frac{r^2 - B^2}{2B\sin(\vartheta - \xi) - 2r} \tag{4}$$
The imaging Geometry is given by
$$\sin(\vartheta - \xi) = \frac{r^2 - B^2}{2BR} + \frac{r}{B} \tag{5}$$
The ground range is determined by
$$y = R\sin\vartheta \tag{6}$$
The local terrain height z above the reference area is therefore:
$$z = H - \frac{r^2 - B^2}{2B\sin(\vartheta - \xi) - 2r}\cos\vartheta \tag{7}$$

I connection with (2) equation (7) delivers the desired Interferogramme as function of the look angle. In connection with (6) equation (7) delivers the ground range representation: The propagation path difference r for the both signals at the two antennas is:

$$r = m\left[\sqrt{R^2 + B^2 - 2BR\sin(\vartheta - \xi)} - R\right] \tag{8}$$

m=1for one pass interferometry, m=2 for two pass interferometry with separate parallel flight passes. The corresponding phase difference Φ is:

$$\Phi = m\,2\pi\frac{r}{\lambda} = \frac{m\,2\pi}{\lambda}\left[\sqrt{R^2 + B^2 - 2BR\sin(\vartheta - \xi)} - R\right] \tag{9}$$

With equations (3) to (9) the topographic height z at each point on the y-axes can be determined as well as all the interrelations between all parameters for one Pixel.

The interferogramme for the whole scene is generated by multiplying pixel by pixel the image I^*_1 with the complex conjugated image I^*_2 of the other. Unfortunately, the phase is represented as modulo 2π only, the phase is wrapped. Thus a "phase unwrapping" has to be carried out to obtain the absolute phase, which finally can be transformed to the desired height z at each point over the x-y-area leading to a three dimensional terrain elevation model.

For simplification it will be assumed B<<R which is realistic for the most airborne and all space-borne SAR systems used up to now. Under that condition results from a Taylor development of (9)

$$\Phi \approx m\frac{2\pi}{\lambda}\left[\frac{B^2}{2R} - B\sin(\vartheta - \xi)\right] \tag{10}$$

For both space-borne and the most airborne SAR the first term can be neglected also because

$$\frac{B^2}{2R} << B\sin(\vartheta - \xi)$$

For the most operational SAR Systems results:

$$\Phi \approx -m\frac{2\pi}{\lambda}B\sin(\vartheta - \xi) = -m\frac{2\pi}{\lambda}B_{par} \tag{11}$$

In (11) the factor $B_{par} = B\sin(\vartheta - \xi)$ is the Baseline component parallel to the look direction.

2.2 Interferometric Performance, Sensitivities and Errors

All equations content both Radar System parameter and Interferometry parameter. For basic considerations the following Radar System parameter, principally, are assumed to be known a priori: The height H above a reference area (a special Geoid ore a flat area respectively), the baseline B, the wavelength λ used, the platform attitude represented by the baseline tilt angle ξ, the look angle ϑ, and the slant range distance R between the pixel observed at P_0 the both antennas A_1 and A_2 and the systems Phase Noise.

However, in reality all these parameters have to be determined by measurements during the flight and, therefore, errors will be introduced due to unavoidable measurement inaccuracies. That will be demonstrated exemplary for the dependency of the elevation variation Δz from the height variation ΔH. Principally holds under the assumption of small pixels for the variation of z as a function of the variation ΔH of H:

$$\frac{\Delta z}{\Delta H} \approx \frac{\partial z}{\partial H} \text{ and therefore } \Delta z \approx \frac{\partial z}{\partial H}\Delta H \tag{12}$$

The exemplary starting equation is (7) which gives the elevation z of the point P in Fig.2. In connection with (7) results for the elevation sensitivity Δz:

$$\Delta z = \Delta H \tag{13}$$

From the sensitivity Δz the quadratic error σ_z results in that case:

$$\sigma_z^2 = \left(\frac{\partial z}{\partial H}\right)^2 \sigma_H^2$$

The entire elevation measurement accuracy depends on the accuracy of the determination of both all the Radar System parameters and all the Interferometry parameters.

The general expression for the elevation errors is derived by adding the all quadratic errors:

$$\sigma^2_z = \left[\frac{\partial z}{\partial H}\right]^2 \sigma^2_H + \left[\frac{\partial z}{\partial R}\right]^2 \sigma^2_R + \left[\frac{\partial z}{\partial \xi}\right]^2 \sigma^2_\xi + \left[\frac{\partial z}{\partial \lambda}\right]^2 \sigma^2_\lambda + \left[\frac{\partial z}{\partial \Phi}\right]^2 \sigma^2_\Phi + \left[\frac{\partial z}{\partial B}\right]^2 \sigma^2_B \quad (14)$$

The subscripts in (14) determine the respective interferometry parameter.

Special interest has to be addressed to the 5th term, the Phase errors. The total phase results from the sum of several contributions. The designations are self explaining:

$$\Phi = \Phi_{flatearth} + \Phi_{topography} + \Phi_{atmosphere} + \Phi_{motion} + \Phi_{noise} + \Phi_{Decorrelation} \quad (15)$$

The sensitivity of a function $\Phi = \Phi(z,y)$ against variations Δz and Δy respectively results from differentiation of the appropriate equations and can be expressed by the total Differential

$$d\Phi(z,y) = \frac{\partial \Phi}{\partial z}dz + \frac{\partial \Phi}{\partial y}dy \quad (16)$$

By replacing the symbol d with Δ and using the equations (7), (9), and (12) results:

$$\Delta\Phi = \frac{\partial \Phi}{\partial z}\Delta z + \frac{\partial \Phi}{\partial y}\Delta y \text{ with } \frac{\partial \Phi}{\partial z} = \frac{\partial \Phi}{\partial \vartheta}\frac{\partial \vartheta}{\partial z} \text{ and } \frac{\partial \Phi}{\partial y} = \frac{\partial \Phi}{\partial \vartheta}\frac{\partial \vartheta}{\partial y}$$

From (11) follows :

$$\frac{\partial \Phi}{\partial \vartheta} = -m\frac{2\pi}{\lambda}B\cos(\vartheta-\xi) = -m\frac{2\pi}{\lambda}B_{eff}$$

The factor $B\cos(\vartheta-\xi)$ is called the effective Baseline B_{eff} vertical to the looking direction.

$$B_{eff} = B\cos(\vartheta-\xi) \quad (17)$$

With (6) follows both:

$$\frac{\partial \vartheta}{\partial z} = \frac{1}{R\sin\vartheta} \text{ and } \frac{\partial \vartheta}{\partial y} = \frac{1}{R\cos\vartheta} \quad (18a) \text{ and } (18b)$$

Therefore, for the total phase variation in z- and y-direction results:

$$\Delta\Phi = -m\frac{2\pi}{\lambda}\frac{B}{R}\frac{\cos(\vartheta-\xi)}{\sin\vartheta}\Delta z - m\frac{2\pi}{\lambda}\frac{B}{R}\frac{\cos(\vartheta-\xi)}{\cos\vartheta}\Delta y \quad (19)$$

$$-\pi \leq \Phi \leq \pi \mod 2\pi$$

This equation, principally, answers the question: Which Phase variation $\Delta\Phi$ will be caused through terrain variations Δz in vertical ore Δy in horizontal direction. It delivers the interconnection between the measured phases and the height characteristics of the observed pixel. The first term represents the phase sensitivity against elevation variations in z-direction, the second term represents the sensitivity against a horizontal shift in a constant altitude z in y-direction taking into account the relations (3) and (6) again. Both terms represent interference patterns.

The second term is caused from the change in slant range observation distance R over the whole swath causing a fringe pattern parallel to the flight direction also. Equation (19) shows: the flat earth causes a fringe pattern which overlays the topographic fringes additionally and, in order to observe and to analyze the topographic fringes only these two effects have to be separated by subtraction of the flat earth interferogram from the total observed interferogrametric phase (19).

The flat earth pattern is always present within the measurements because both R and (R + r) increase with increasing incidence angle, and B_{eff} , principally, changes its value also. The flat earth pattern has to be removed in order to obtain the pure elevation above the reference area which in this case is the flat area at z=0.

$$\Delta\Phi_z = -m\frac{2\pi}{\lambda}\frac{B_{eff}}{R\sin\vartheta}\Delta z = -m\frac{2\pi}{\lambda}B_{eff}\Delta\vartheta_{vert} \quad (20)$$

$$\Delta\Phi_{flat} = -m \frac{2\pi}{\lambda} \frac{B_{eff}}{R\cos\vartheta} \Delta y = -m \frac{2\pi}{\lambda} B_{eff} \Delta\vartheta_{hor} \tag{21}$$

The relation (20) delivers the required interference pattern between the measured phases and the terrain elevation z, i.e. a vertical shift of the observed point P, the relation (20) delivers the interference pattern which results from a horizontal shift of the observed pixel. $\Delta\vartheta_{vert}$ and $\Delta\vartheta_{hor}$ are the small looking angle variations for the horizontal and the vertical shifts of the point P respectively. From the geometry depicted in Fig.3 the following relations for $\Delta\vartheta_{vert}$ and $\Delta\vartheta_{hor}$ can be determined:

For the horizontal shift:

$$R\Delta\vartheta_{hor} \approx R\sin(\Delta\vartheta_{hor}) = \frac{\Delta R}{\tan\vartheta} \tag{22}$$

For the vertical shift:

$$R\Delta\vartheta_{vert} \approx R\sin(\Delta\vartheta_{vert}) = \frac{\Delta z}{\tan\vartheta} \tag{23}$$

Here with result from (19) and (20) the phase sensitivities for elevation and slant range variations.

$$\Delta\Phi_{flat} = -m \frac{2\pi}{\lambda} \frac{B_{eff}}{R\tan\vartheta} \Delta R \tag{24}$$

$$\Delta\Phi z = -m \frac{2\pi}{\lambda} \frac{B_{eff}}{R\sin\vartheta} \Delta z \tag{25}$$

$$-\pi \leq \Phi \leq \pi$$

The equations (21) and (24) represent the interferogram of the flat earth, i.e. an interferogram in absence of any topographic changes. The equations (20) and (25) represent the so called "flattened topographic interferogram" of the observed area in slant range representation. Both inerferogrammes have to be transformed into ground range presentations in order to obtain the desired results. The interconnection between ground range R_g and slant range R is:

$$R_g = R\sin\vartheta$$

With that relation results for the both Interferogrammes:

$$\Delta\Phi_{flat} = -m\frac{2\pi}{\lambda}B_{eff}\frac{\Delta R}{R_g}\cos\vartheta \tag{26}$$

$$\Delta\Phi_z = -m\frac{2\pi}{\lambda}B_{eff}\frac{\Delta z}{R_g} \tag{27}$$

In all equations varies Φ mod 2π between -π and +π, whilst $\Delta\Phi$ varies between -π and +π only. The width of the fringes, principally, is given for $\Delta\Phi = 2\pi$. The sensitivity depends on the minimum phase measurement capability of the sensor. The state of the art is $\Phi \approx 5°$ which roughly corresponds to 1/70 of a wavelength. Fig.4 shows modeled fringe patterns. In Fig. 5 a measurement result is depicted, in the figures 6, and 7 for an space-borne and an airborne case measured interferogrammes over partially relative flat areas are depicted.

For the space-borne case with its large altitudes and small incidence angle variations a nearly periodical fringe pattern results with nearly equal fringe width's result. The flat-earth phase is a linear phase ramp.

For airborne data the situation is more complicated, the look-angle changes strongly from near-range to far-range, and with it both the observation distance r and the effective base line. Therefore nearly everything gets range-dependent. The flat-earth phase is not a linear phase ramp anymore, it has a curved shape. The height resolution changes with the incidence angle also.

2.3 The Spectral Approach to Interferometry

The relation between the phase of SAR interferograms and ground elevation is explained as usual by means of geometrical approaches based on the assumption that the RF bandwidth is so sma to be negligible. Thus, the system is considered monochromatic. However, if the finite system bandwidth is considered, a relative shift of the ground wave number spectra dependent on the baseline and the local slope is recognized. Important consequences are revealed from this result. [Gatelli et al. 1994]

1) It can be shown that as the baseline increases and in the absence of volumetric effects the spatial resolution of the interferogram is reduced only since the "geometrical de-correlation" can be removed [Prati C., Rocca F., 1993]. Therefore, techniques for improving the interferogram quality (i.e., yielding a higher SNR) can be found (with a considerable benefit for phase unwrapping). Moreover, this effect could be completely avoided by means of tunable systems.

2) Rules for the design of pre-summing filters for quick-look interferometric processing can be derived.

3) Repeated satellite surveys can be combined in order to improve the slant range image resolution because the spectra of different signals correspond to different bands of the ground reflectivity's spectrum. The relation between the frequency f and the ground range wave number k_y is given bv:

$$k_y = \frac{4\pi}{\lambda} sin(\vartheta - \alpha) = \frac{4\pi f}{c} sin(\vartheta - \alpha) \qquad (28)$$

Here, for simplicity, a constant uniform slope α of the terrain is considered. For the variation of k_y generated by a slight change of the looking-angle $\Delta\vartheta$ follows:

$$\Delta k_y = \frac{4\pi f \Delta\vartheta}{c} cos(\vartheta - \alpha) \qquad (29)$$

Thus, in general, a looking-angle difference $\Delta\vartheta$ generates a shift and a stretch of the imaged terrain spectra. However, if the relative system bandwidth is small, the frequency f in the second term of (28) can be substituted with the central frequency f_0.

$$\Delta k_y = \frac{4\pi f_0 \Delta\vartheta}{c} cos(\vartheta - \alpha) \qquad (30)$$

From the fact that the radar is not monochromatic, it has a bandwidth W centered around the central frequency f_0 , results that by changing the looking angle of the SAR survey, a different band of the ground reflectivity spectrum shows up.

In order to compare the shift of the ground reflectivity spectrum to the SAR bandwidth W, the ground wave number shift of (30) is expressed as an equivalent frequency shift Δf. The following expression of Δf for an angular separation $\Delta\vartheta$ can be obtained by differentiating (28) directly and taking into account the relations for the small incidence angle changes resulting from Fig. 3:

$$\Delta f = -\frac{f_0 \Delta\vartheta}{tan(\vartheta - \alpha)} = -\frac{cB_{eff}}{R\lambda\, tan(\vartheta - \alpha)} \qquad (31)$$

Note: (31) does not state that by changing the looking angle of the SAR survey the radar bandwidth is shifted by Δf, it just says that by changing the SAR looking angle, the backscattered signal contains different spectral components of the ground reflectivity spectrum. The signals received by two SAR's separated by an angle $\Delta\vartheta$ have the same spectral components of the first signal in the second spectrum shifted by Δf only.

A similar result can be demonstrated for the bistatic SAR case where a stereo pair is acquired simultaneously by two sensors: the first is equipped with a transmitter and a receiver, whereas the second is equipped with a receiver only. It can be shown that the spectral shift induced by a baseline B_{eff} is half that obtained in the monostatic case:

$$\Delta f = -\frac{cB_{eff}}{2R\lambda\, tan(\vartheta - \alpha)} \qquad (32)$$

In the monostatic case the two SAR spectra become totally disjoint as the frequency shift Δf equals the system bandwidth W. The critical baseline can be computed as:

$$B_{effcrit} = \left| \frac{WR\lambda tan(\vartheta - a)}{c} \right| \qquad (33)$$

By using the geometric range resolution δ_{rge} instead of the Bandwidth W results from the well known interconnection between range resolution and bandwidth:

$$B_{effcrit} = \left| \frac{R\lambda tan(\vartheta - a)}{2\delta_{rge}} \right| \qquad (34)$$

The Tables 1 to 3 summarize the results obtained in this chapter.

Ground Range:	$y = R \sin\vartheta$ (6)
Local Terrain Height:	$z = H - \dfrac{r^2 - B^2}{2Bsin(\vartheta - \xi) - 2r} cos\vartheta$ (7)
Critical Baseline	$B_{effcrit} = \left\| \dfrac{WR\lambda tan(\vartheta - \alpha)}{c} \right\| = \left\| \dfrac{R\lambda tan(\vartheta - \alpha)}{2\delta_{rge}} \right\|$ (34)
Surface Change Detection Differential Interferometry	$\Delta z = \dfrac{\lambda}{4\pi} \Delta\Phi$

Tab1 Basic Interferometry Formulas

Input parameters	Value
• velocity: v	7.464 km/s
• range delay (mid swath): T_0	5.779 ms
• slant range (mid swath): R_0	866.656 km
• incidence angle (mid swath): ϑ	23°
• altitude: H	794 km
• wavelength: λ	0.0566 m
• baseline tilt angle: ξ	0°
• baseline length: B	200 m
• surface slope: r_x	r_y
• phase error: $\Delta\Phi$	5°
• baseline length error: ΔB	0.001 m
• baseline tilt angle error: $\Delta\xi$	1°
• slant range error: ΔR	3 m
• height error: ΔH	1 m

Tab. 2 Exemplary Parameters for Interferometry with an ERS/ASAR Type space-borne SAR

Conclusions

Absolute height measurement with interferometric SAR requires an extremely stable system performance. Especially the orbit has to be known very precise due to its great impact on the baseline length, the tilt angle accuracy and the altitude accuracy. An instrument phase error of 5° leads to an error in the height estimation of about 1.5 m in the mid swath and seems to be acceptable.

Error Type	Formula	Errors
Phase	$$\Delta z_\Phi = \frac{\lambda R \sin \vartheta}{4\pi B \cos(\vartheta-\xi)} \Delta\Phi = \frac{\lambda R_g}{4\pi B_{eff}} \Delta\Phi$$	$\Delta\Phi = 5° \rightarrow \Delta z_\Phi$ 0,723m
Baseline Length	$$\Delta z_B = -R \tan(\vartheta - \xi) \sin\vartheta \frac{\Delta B}{B}$$	$\Delta B = 0,001$ m $\rightarrow$, $\Delta z_B = 0,719$ m
Baseline Tilt Angle	$$\Delta z_\xi = -R_0 \sin\vartheta \Delta\xi$$	$\Delta\xi = 1° \rightarrow \Delta z_\xi = 5910,20$ m
Range Delay	$$\Delta z_R = -\cos\vartheta \Delta R$$	$\Delta R = 3$ m $\rightarrow$ $\Delta T = 0,2$ μsec $\rightarrow \Delta z_R = -2,762$ m
Altitude	$$\Delta z_H = \Delta H$$	$\Delta H = 1$ m $\rightarrow \Delta z_H = 1$m

Tab. 3 Resulting Errors for absolute Height Estimation

DEM Error Examples for SRTM (Middle of Swath)				
	Relative (30 seconds)		Absolute (11-days)	
Error Type	Inaccuracy	DEM-Error	Inaccuracy	DEM-Error
Baseline Tilt Angle	2 arcsec	3,0 m	9 arcsec	13,4 m
Baseline Length	1,3 mm	0,8 m	4,0 mm	2,6 m
Instrument Phase	4,0deg	4,2 m	4,0 deg	4,2 m
Total (RSS)	5.5 m		14.4 m	

Tab.3 Digital Elevation Model errors caused by instrument inaccuracies of SRTM/X-SAR. Goal was a relative height accuracy (90 %) < 6 m absolute height accuracy (90 %) < 16 m [Keydel et al 2000]

3. Coherence and Correlation in SAR Interferometry

For SAR Interferometry the coherence between two signals to be compared defined in the Introduction Lecture is a fundamental parameter. The Coherence is a measure for the quality of interferogrammes. Equation (4) will be used to produce the complex SAR-Interferogram replacing the Fields E_1 and E_2 by the complex Images I_1 and I_2 gained in each of the two interferometry channels. This interferogram is defined as the product of the complex SAR values of the second image with the complex conjugate of the reference image. Principally, the Interferogram is the complex Coherence Image C_{image}:

$$C_{image} = \frac{\left\langle (I_1(t)I_2^*(t)) \right\rangle}{\sqrt{\left\langle I_1(t)I_1^*(t) \right\rangle \left\langle I_2(t)I_2^*(t) \right\rangle}} \tag{28}$$

The corresponding amplitudes have to be averaged and the corresponding phases have to be differenced at each point in the image. The resultant phase of the interferogram, which is directly related to the topography, is measured modulo 2π only. To calculate the elevation of each point it is, therefore, necessary to add the correct integer number of phase cycles to each phase measurement. The problem of solving this ambiguity is called phase unwrapping.

On that basis applications of interferometry can be split into two categories, based on whether the main parameter of interest in the complex interferograme is the phase (4) or the magnitude (5) of the complex coherence.Fig 8.

The phase information can be used for digital elevation model (DEM) production, provided the scene is relatively stable. For repeat-pass interferometry, the phase may also be used for studying small coherent motions of the target between the imaging times (a motion in which most of the scatterers move simultaneously). This has applications in geological studies of land motion (e.g. associated with seismic activity or land slip), and in glaciology. The accuracy of the measurements of height and motion are in part determined by the noise of the phase measurements. This phase noise is usually characterized by the magnitude of the coherence of the interferogram (often referred to as simply the coherence).

Important information can be extracted from images using the coherence. This is true for land use classification, where the information obtained is complementary to the intensity of the SAR images, and may give, for example, better discrimination between clear cuts and forest areas than intensity measurements. Another use of coherence in this respect is the monitoring of changes to land surfaces (e.g. harvesting of agricultural crops). Fig 11 shows an example for Land classification obtained with coherence evaluation of a multi frequency X-SAR/SIRC result. The respective phase image without flat earth removal is depicted in Fig..

There is a number of causes for de-correlation in interferometry, and it is possible to write the coherence as a product of factors which give some physical insight into the origin of the decorrelation. It is possible to write the coherence as a product

$$\gamma = \gamma_{instrument}\, \gamma_{processor}\gamma_{baseline}\, \gamma_{azimuth}\, \gamma_{noise}\, \gamma_{temporal}\, \gamma_{spatial}$$

The first three terms are totally independent of the target being imaged, the fourth and fifth are related to both the system and the target, and the final term is totally dependent on the target properties as will be explained.

$\gamma_{instrument}$: refers to both Radar and Platform

$\gamma_{processor}$: refers to effects produced during the processing of the SAR images to form the interferogram and includes, for example, miss-registration of the images when aligning them before forming the interferogram and overlap of the range and azimuth spectra. The latter means the two surveys must image the scene under the same squint angle, they have to have the same Doppler Centroid. The to days processing routines are estimated to register the two images to within one hundredth of a pixel in both range and azimuth, leading to a negligible loss of coherence.

$\gamma_{azimuth}$ is a factor introduced to account for differences in the Doppler centroid frequencies between the two images. For repeat-pass interferometry with the ERS-1 satellite this effect is negligible as the high stability means that the Doppler centroids typically vary by only about 0.02 of the sampling frequency.

γ_{noise} represents the effect of noise in the system is well known and can be expressed as

$$\gamma_{noise} = \frac{CNR}{CNR+1} = \frac{1}{1+CNR}$$

where CNR denotes the clutter-to-noise ratio. This is important for regions of weak backscattering.

$\gamma_{temporal}$: the decorrelation in single pass interferometry would depend simply on the factors discussed above. In the case of repeat pass interferometry, where the time interval between images may be days or longer, there is also a contribution to the decorrelation due to the possibility of changes in the scatterers themselves and due to the changes of weather conditions in the propagation path. An agricultural field for example de-correlates steadily with the evolution of the crop. The correlation is completely destroyed once the field is ploughed. Forests de-correlate even more rapidly etc. On the other hand many manmade and

natural objects prove to stay coherent over years and decades which enables the techniques of Permanent Scattering Interferometry considered in Chap.7

$\gamma_{baseline}$: the baseline of the two passes has to be smaller than the critical effective baseline

$\gamma_{spatial}= \gamma_{range} \gamma_{volume}$ Caused by different acquisition look angles

γ_{range} : Decorrelation due to different projections of the bandwidth onto ground

γ_{volume}: Decorrelation related to the height distribution of the scatterers

In Fig. 8 the influence of coherence degradations on the interferogram is depicted. The definition of coherence is impossible to apply directly to the SAR images, because of the fact that it involves ensemble averaging. In practice an assumption is made in which a spatial average is taken to be a reasonable approximation of an ensemble average. This requires that phase variations between the pixels being averaged have been corrected for topographic variations. The averaging process introduces a bias into the coherence measurements, the level of which is determined by the number of independent samples used.

However, the coherence of a pair of interferometric images gives one indication of the quality of the obtained interferogram [Monti97]. Also the coherence matrix is a basic starting point for some unwrapping algorithms. If real time is desired, or an increment in the speed of all the chain from raw data to the digital elevation model (DEM) is needed, the time reduction in computing the coherence should be undertaken.

The usual coherence estimator needs to calculate the local frequency of the interferometric fringes very accurately and therefore this algorithm becomes very time consuming. The coherence between two complex SAR images I_1 and I_2 is given by (28). Assuming stationary regions and ergodicity (i.e. all scatterers within the estimation area are independent with identical statistical properties), the coherence can be calculated with high accuracy as:

$$\gamma_{est}=\frac{\left| \sum_{n=0}^{N} \sum_{m=0}^{M} I_{1(n,m)} I_{2(n,m)} e^{-j\Phi(n,m)} \right|}{\sqrt{\sum_{n=0}^{N} \sum_{m=0}^{M} \left| I_1(n,m) \right|^2 \sum_{n=0}^{N} \sum_{m=0}^{M} \left| I_2(n,m) \right|^2}} \qquad (29)$$

γ_{est} is called coherence estimator. N and M give the size of the estimation window in range and azimuth. The factor The inclusion of a factor of $e^{-j\Phi(n,m)}$ is due to the difference between the two SAR images, which differ by the deterministic phase, that should be estimated and corrected.

4. Generation of Interferometric Products

The input data used for SAR interferometry are the complex SAR image pairs and the respective navigation data. The procedure used for the generation of the interferometric products is depicted in Fig.12 and can be summarized as follows:

• Range and azimuth filtering of the complex images. Before the interferogram is formed, the complex image pair has to be separately filtered in both range and azimuth directions to optimize the coherence and the interferometric phase purity. The so-called spectral shift range filtering was applied in order to minimize baseline-induced decorrelation. Filtering to a common Doppler band in the azimuth direction was applied to eliminate the decorrelation effect induced by the Doppler centroid difference between the two images [J. Moreira, M. Schwäbisch,et.al 1995].

• Coregistration. In order to obtain the interferometric phase, the image pair has to be co-registered. This has to bo be carried out with a high degree of accuracy state of the art is the order of one hundreth of a pixel.

• Interferogram formation. The interferogram is generated by multiplying the first image, L , with the complex conjugate of the second image, referred to as I_2

• Flat-earth phase removal. To perform the interferogram noise filtering and simplify the phase unwrapping procedure, a first-order, flat-earth phase removal is carried out. The flat-earth phase is computed by

measuring the dominant fringe frequency in the azimuth and range directions of the interferogram. It thus consists of two different linear phases in the range and azimuth directions..

• Phase noise reduction. In order to reduce the phase noise and to facilitate the phase unwrapping, the interferogram has to be filtered using an averaging window with a size of several resolution elements in both range and azimuth. Respective pixel spacings and resolutions used for X-SAR processing are shown in Table 1 exemplary. After the filtering, the interferometric phase is determined by calculating the arctangent of the interferogram. The terminology and respective procedures are given in the next sections.

4.1 Phase Unwrapping

Phase Unwrapping is the key procedure for SAR interferometry data evaluation with respect to establishing DEM's. Meanwhile, there are many methods for the two dimensional Phase unwrapping.

The very basic principle is depicted in Fig.11 for a simple one dimensional case which is self explaining. A more sophisticated principle is the least square-Method which minimizes the differences between the derivatives of the wrapped and unwrapped phase functions using the fact that the minimum of the first derivative is determined by the null of the second derivative.. In the two dimensional case the partial differential quotients have to be used. This method seems to be very simple and fast. However, the errors are propagating very strongly down to the entire sequence and therefore, this method is not very practicable for large area processing.

As further examples: branch cutting, fringe detection, and knowledge injection should be mentioned in a bit more detail despite of the fact that they don't represent the most modern state of the art. However, they seem to be important to gain insight into the whole complex. [Gens R., 1996]

Branch cuts

The first method, which is based on creating so called *branch cuts* was developed in 1988 already by Goldstein *el al.* They pointed out that there are two possible types of errors in the unwrapped sequence. The first are local errors in which only a few points are corrupted by noise; the second are global errors in which the local error may be propagated down the entire sequence. The global errors arise from the residues. Residues are local errors in the measured phase caused by the noise in the signals or by actual discontinuities, i.e., layovers, in the data. There are three major sources of noise. The first is thermal noise, the second is the speckle, which is a random variable, and the third one is caused by an antenna effect depending on the antenna separation, the bandwidth of the radar, and the geometry of the image formation. Assuming that the original scene is sampled often enough so that the true phase will not change by as much as one-half cycle per sample point the phases can be unwrapped in a consistent manner. This can be done if the residues are identified and suitable branch cuts are made between the residues to prevent any integration path from crossing these cuts. Filtering the signal will reduce the residues in the image but it also reduces the spatial resolution in the final topographic map. Before integrating the phase differences all residues which were previously low-pass filtered are to connect with branch cuts. The phase differences are then integrated without crossing any of the cuts.

Fringe detection

Another method for the two-dimensional phase unwrapping is the fringe detection method of Lin *et al.* (1992). The first step in this method is to find the location of the fringe lines in the phase difference image by using edge detection techniques which mostly include the two steps of enhancement by filtering and then applying a threshold to the image. The phase difference image is median filtered with a window size which depends on the density of the fringes. For the detection of the fringe lines, five edge masks with different orientations are applied. The phase is then unwrapped by adding a multiple of 2π each time the integration path crosses a fringe line. The multiple depends on how many fringe lines lie between the pixel and a reference pixel. The solution is considered good when the fringes are greatly separated and the signal-to-noise ratio (SNR) is high in the entire wrapped phase image [Hartl et al.1993].

Knowledge injection

This method is used for reducing the density of fringes by incorporating the use of a rough DEM. [Adragna 1995]. The same technique is also used for flat Earth removal, and for differential interferometry (Cap.5.),

in which knowledge of the topography is injected to isolate the terrain movement. It should be noted that all the methods of phase unwrapping described above are generally performed after flat earth removal.

4.2 Co-Registration

The co-registration is a very important processing step in interferometry also. For distributed scatterers, a co-registration inaccuracy introduces substantial noise into the interferometric phase (Bamler et al., 1993), and point scatterers have to deal with a systematic phase error (Holzner et al.,2001). (Compare the PS considerations in Chap.8 also). Bamler (2000) showed that the error of the co-registration parameter estimation for distributed scatterers depends on the number of correlation points n and the coherence γ:

$$\sigma_{corr} = \sqrt{\frac{3}{2n}} \frac{\sqrt{1-\gamma^2}}{\pi\gamma}$$

The dimension here is the unit of the resolution cell. Accordingly, in order to guarantee an appropriate co-registration accuracy in low coherence scenes the correlation window size needs to be increased. However, the co-registration still strongly depends on the temporal coherence and the co-registration limit is often insufficient.

5. Along-Track Interferometry

The arrangement of the two antennas in cross velocity direction is essential for topographic mapping because that arrangement only leads to phase differences between the two SAR images. An along track arrangement as depicted in Fig.2 would produce two identical images with no phase differences at all due to the same processing procedure. Each object had both the same distance and the same Doppler Frequency with respect to the two receivers ore antennas respectively. However, there is a time shift T between the two images due to the fact that both antennas pass the point of nearest approach to each pixel - which is the locus of Doppler Frequency 0 - in a sequence which is dependent on Baseline length and platform velocity v_{plat}.

$$T=B/v_{plat}$$

A target changing its position during that time with the velocity v_{targ} will change its position in the second image in accordance with its own velocity producing a phase difference Φ. The Phase difference between corresponding pixels is a measure for the velocity of the object in the scene. This procedure is called Along Track Interferometry.

$$\Phi = m\frac{2\pi}{\lambda}v_{targ}T = m\frac{2\pi}{\lambda}\frac{v_{targ}}{v_{plat}}B \qquad (30)$$

In order to determine the ground range velocity of a target ore a pixel resp. moving on ground neglecting the vertical velocity component $v_{targ\,z}$ for simplicity reasons results:

$$V_{targ} = \frac{V_{targ}}{\sin\vartheta} \qquad (31)$$

The measurement of the velocities of different resolution cells of water surfaces for instance allows the estimation of water currents with respect to both current velocity and the respective direction. However, the along-track technique will not be suitable for the measurement of along track velocity components. But with two separate measurements in cross directions a two dimensional current map can be established. Fig. 12 shows an example. In Fig. 13 results of moving target detection and measurements are depicted.

6. Differential SAR Interferometry

Many operational monitoring tasks require the detection of changes during a certain time period. This can be performed with Differential SAR Interferometry which registers the respective topographic changes of a

DEM which happened during the time period since the DEM has been established. There are principally two techniques depending on the availability of pre-existing digital topographic maps.

If such a DEM with acceptable accuracy is available then it has to be subtracted from the DEM which results from the interferometric phase field measurements. The difference between the two DEM's represents the changes to be determined.

If no digital topographic maps are available then two interferograms of the same area have to be taken over the desired time period. The first one is used to generate a reference Interferogram and the second one to generate a Interferogram which contents both the reference DEM together wit the changes to be monitored. This leads to a double-difference interferogram, which can be derived from the two corresponding interferograms by subtraction of both interferograms.The phase discontinuities in the two interferograms will not occur in the same places. Therefore, each interferogram must be unwrapped.

The sensitivity of SAR interferometry depends on the geometry of the observation and can be increased at the expense of spatial resolution by averaging over more pixels. The length of the baseline in relation to the wavelength used has an influence on the amount of noise which is introduced. A decreasing baseline ore an increasing wavelength leads, therefore, to a higher accuracy of the detected height differences.

The principal geometry for the space borne two pass mode is depicted in Fig 14. For simplicity reasons in describing the principles it is assumed that both Interferograms use the same primary flight path, however they can have different baselines. The following Interferograms will be obtained:

Interferogram 1:

$$\Phi_1 = m\frac{2\pi}{\lambda}r_1 = \Phi_{topography} = m\frac{2\pi}{\lambda}B_{1\,par} \qquad (32)$$

Interferogram 2:

$$\Phi_2 = m\frac{2\pi}{\lambda}r_2 + m\frac{2\pi}{\lambda}\rho = \Phi_{topography} + \Phi_{displacement} \qquad (33)$$

The subtraction of (32) from (33) leads to the displacement in line of sight direction

$$\Phi_2 - \Phi_1 = \Phi_{displacement} = m\frac{2\pi}{\lambda}\rho \qquad (34)$$

This expression is not dependent on the topography anymore. As explained in Cap. The sensitivity results from differentiation which leads to

$$\Delta(\Phi_2 - \Phi_1) = \frac{2\pi}{\lambda}\frac{\Delta z}{\cos\vartheta} + \frac{\Delta y}{\sin\vartheta} \qquad (35)$$

Fig 15 shows exemplary the displacements caused by an earthquake in Japan.

7. Permanent Scatterer Interferometry

Differential Interferometric SAR over longer time periods is often limited by temporal de-correlation. As mentioned already many manmade and natural objects prove to stay coherent over years and decades. These so-called Permanent Scatterers (PS) can be used for the ultimate sensitivity in motion measurement from space (Ferretti et al., 2000; Ferretti et al., 2001) by using them as points of reference. PS's are usually bright points in the image, with a very high signal-to-noise ratio. However, the phase information itself cannot be used for detection of stable points, since it contains phase uncertainties induced by unknown elevation and deformation but amplitude time series can be well used for this [Ferretti 2001]. The permanent Scatterer technique uses a sequence of many Interferometric SAR products of the same area of interest taken over a long time period establishing in a first step a data base of very many SAR images. From the observed phase time series of the differential interferograms the elevation difference and deformation rate difference over time can be determined.

Presently such databases are available, principally, from the most areas of the globe taken from different operational satellites like ERS-1/2, RadarSAT, J-ERS etc. which where in orbit permanently since more than 15 years. (Some of them will be numbered and considered in the Systems chapter of this very lecture Series).

If several SAR images are superposed, even the propagation errors due to tropospheric water vapour can be averaged out. In practice 30 – 100 SAR images of the same areas are used to form an interferometric stack spanning a time period of typically 5 – 12 years. In this co-registered stack, the PS's are identified and their phase history can be tracked and analysed. Several filtering and adjustment steps are required to remove orbit errors, long-wavelength ionospheric delays and to reduce atmospheric effects and noise. The result is a map of subsidence/inflation rates to an accuracy of less than 1 mm/year.(Bamler, SRTM and beyond)

If, for example, n + 1 ERS SAR images from one certain area are available the data will be first co-registered on a unique master and a first DEM of the area will be established as a reference. Then, n differential interferograms between all SAR images and the master will be computed. After DEM compensation, the residual phase Φ_i of the interferogram number i is:

$$\Phi_i = \frac{4\pi}{\lambda} r_i(T_i) + \Phi_i^{atm} + \Phi_i^{noise} + \Phi_i^{topo} + \Phi_i^{displacement} + \Phi_i^{orbit} \quad i = 1..n \quad (36)$$

$r_i(T_i)$ represents the momentary target position at the time T_i and herewith, principally, the possible target motion in the line-of-sight direction, Φ_i^{atm} is the atmospheric phase contribution, Φ_i^{decorr} the contribution of the de-correlation noise, Φ_i^{topo} is the topographic phase error etc. All errors are mentioned already in Chap.3.4,

A time series analysis of the first term in equation 1 reflects the target motion: here for simplicity a constant velocity model is supposed. An uniform strain rate hypothesis is often used in geophysical modelling; however, more complex models can be adopted if the linear model is unrealistic. Permanent scatterers are characterized by low phase residues and are determined by

$$\frac{\partial}{\partial T}\left(\frac{4\pi}{\lambda} r_i(T_i)\right) = 0$$

The first term in equation 1 can be written as follows:

$$\frac{4\pi}{\lambda} r_i = \frac{4\pi}{\lambda} V_r T_i = C_{Vi} V_r \quad i=1............n \quad (37)$$

V_r is the unknown component of the target velocity m in the direction of LOS and T_i is the temporal baseline between the master acquisition and the generic i-th slave image. Since there are n differential interferograms of the same area with different temporal and geometric baselines, a linear system of n equations and, principally 2 unknowns (the velocity V_r and the DEM-quantities) is established and a joint estimation of both DEM errors and target velocity can be carried out on a pixel-by-pixel basis. In order to reduce the effect of atmospheric patterns and compensate for baseline errors, the low frequency components should be previously removed from Φ_i. The processing is then performed on a pixel-by-pixel basis and the final result is a velocity field of the area of interest together with an improved DEM. In order to compensate for possible baseline errors and/or very low frequency spatial phase distortion due to atmospheric inhomogenities the linear component of the residual phase values was removed from all the available data.

Fig.16 shows the structure of an overall PSI system. (Adam, et. al. 2003). The meaning and the tasks of the single Boxes are as follows:

The Data Import box starts the initial data setup and selects the area of interest. The InSAR processing box represents the normal interferometry processing procedure. The master scene is selected in order to start the InSAR Processing. Parameters as the effective baseline, the acquisition date, the Doppler centroid frequency and the season of the acquisition form the selection criteria. During the InSAR processing the observation geometry parameters as the height-to-phase conversion factor, the flat-earth phase, the range distances and the look angle are computed. Furthermore, the scenes are co-registered and re-sampled with respect to the master scene and the interferograms are generated. Because of the utilisation of point scatterers, a spectral shift filtering is not necessary.

In the Differential Interferometric Processing module the observation geometry of the radar acquisition is simulated. A DEM and precise orbits are used as input in order to correct for the interferograms. Any available DEM can be used to reduce the topographic signal. Exemplary the DEM's generated from the X-SAR shuttle mission can be used ore a DEM derived from the ERS tandem mission. This module also determines the co-registration and, therefore, it can be used for that processing step also. With the PS interferometry, the co-registration becomes more difficult than in normal Interferometry considered in Chap.4. The reason is the long time separation (up to ten years) between the acquisitions and the immense temporal decorrelation in non-urban areas.

The calibration of the scenes has to be introduced for the PS detection. To analyse the temporal backscattering behaviour of the point scatterers it is necessary to compare their intensity by the calculation of the radar cross section. This module corrects for the processor gain constant, the antenna pattern and the range spreading loss. The Permanent Scatterer Detection and identification is performed in the co-registered calibrated scenes by a homogeneity test (Ferretti et al., 2001). Aim is to find as many scatterers as possible because a subsidence pattern and the atmosphere have to be sampled spatially as dense as possible. However, unreliable points causing incorrect estimation have to be avoided. Therefore, the temporal analysis of the differential phases is restricted to point scatterers with a high SNR and a long-time stable backscattering behaviour. These are usually man-made features (Usai et al., 1999) and have to be detected in the scene. The extraction of the PS information reduces the amount of data from more than 100 GB to less than 300 MB. Range and azimuth position of the scatterers in the scene, the calibrated intensity , the look angle, the Doppler Centroid frequency, the height-to-phase conversion factor, the temporal baseline, the scatterer sub pixel position and the differential phase in a compact form are the result from this processing step which can be considered as a raster to vector converter.

The modules Estimation and Visualisation: The measured differential phase is composed of contributions from the uncompensated topography Because the phase is measured modulo 2π a non-linear inversion problem is given. It is solved by utilising the different behaviour of the contributions regarding the acquisition parameters effective baseline, temporal baseline, range and azimuth location of the scatterers. Figure 2 visualises the different spatial properties for a subsidence, for an orbit error, for the atmosphere and for noise. A relative estimation between point scatterers located closely to each other reduces the influence of the atmosphere and of orbit errors. Utilising a periodogram a constant relative subsidence rate can be estimated. These relative estimates are transformed into a global subsidence map by a 2-D integration procedure based on a LS-adjustment.

Fig.17 shows a section of a radar image of Berlin. The PS's are marked by coloured dots, the colours represent the subsidence rate. A second example is shown in Fig.18. The area around the Olympic stadium in Berlin is subject to ground water regulation activities. This area shows up as red PS's, i.e., the ground level is lifting by about 3 mm/year. During the lecture on applications additional examples will be discussed.

8. Multifrequency SAR Interferometry

A multi-frequency SAR system offers a great potential to improve the unwrapping performance of interferograms. The SIR-C/X-SAR Shuttle Mission flown twice in April and October 1994 provided for the first time two pass interferometric data in X-, C-, and L-band gained simultaneously over some test areas. Respective investigations have shown that the availability of multi frequency information permits a drastic improvement of the phase unwrapping performance investigation [Moreira J. et. al 1995]. In the Fig 19 to 21

the sets of tree co-registered slant range images, coherence images, and a set of slant range interferometric phases overlapped with the interferometric magnitude images in L-, C-, and X-Band is depicted.

Considering an image pair which meets the baseline and Doppler centroid requirements, the interferometric phase quality is limited by the temporal decorrelation as has been mentioned already. The height accuracy depends on both the wavelength and the coherence. For areas with the same coherence for L-, C-, and X-band, the height derived from the X-band phase is approximately two times more accurate than C-band and eight times more than L-band due to the wavelength ratios.

In the example of SIR-C/X-SAR, the basic shape of the fusioned DEM, i.e. the low frequency part, is mainly derived from L-band due to its high coherence. The high frequency part of the DEM is derived from C- and X-band due to their small wavelengths. This fusion is done by using a simple Kalman filter and takes into account the respective coherences.

The phase gradient estimation can be drastically improved in the presence of critical areas, i.e. low-coherence areas, if an a priori knowledge of the terrain topography is available. The main idea is to improve the unwrapping procedure in C- and X-band by using, as a priori information, the L-band unwrapped values. The selection of the L-band data is related to its higher coherence and the lower sensitivity, compared to C- and X-band data, of the interferometric phase to the variations of the height profile.

A block diagram of the multi frequency unwrapping procedure is depicted in Fig 22. The first step of the proposed procedure is represented by the filtering operation of the L-band wrapped phase (via a 3 x 3 pixel pivoting median filter) followed by the phase unwrapping step. The L-band unwrapped data are used to improve the C-and X-band unwrapping results. Following the procedure depicted in Fig.22.

First of all, the resulting L-band phase is scaled with respect to the C-band wavelength and subtracted from the C-band measured phase. This operation is carried out in the complex domain so that:
• both rescaled L-band and C-band phases are transformed into complex signals by using an unitary amplitude,
• the difference operation is realized by multiplying the complex C-band signal by the complex conjugate of the L-band signal, and
• the wrapped phase difference is obtained from the arctangent of the above-mentioned complex multiplication
.

The result of the difference operation is filtered with a 3 x 3 pivoting median filter and then unwrapped. We explicitly note that the pivoting median filter will benefit from the improved phase field stationarity due to the performed subtraction. Finally, the computed phase is added to the L-band reference signal.

The same strategy is used to unwrap the X-band phase interferogram. The only difference in this case is that the previously computed C-band phase, instead of the L-band one, is used as a reference pattern. The overall procedure block diagram is depicted in Fig. 23.

To demonstrate the improvement achieved by using the multi-frequency information the corresponding results obtained by independently unwrapping the three interferograms without any filtering and filtered are depicted in Fig.24 and Fig. 25 which show a slices through an unwrapped Interferogam before and after filtering. The comparison makes evident the joint effects of the filtering step combined with the introduction of the multi-frequency information. In Fig. 24 there are still some reconstruction problems caused by extremely low-coherence areas present in the X-band data where the bias could not be completely avoided. A better result can be obtained by applying stronger data filtering. This is shown in Fig. 25, where the L-band reference is compared to a highly filtered X-band unwrapped pattern (retrieved in the multi-frequency phase unwrapping procedure by using the cascade of two 5x5 median filters). Obviously, the improvement obtained is at the expense of a loss in the height resolution in X-band. In Fig. 26 is the resulting DEM of the mount Etna depicted.

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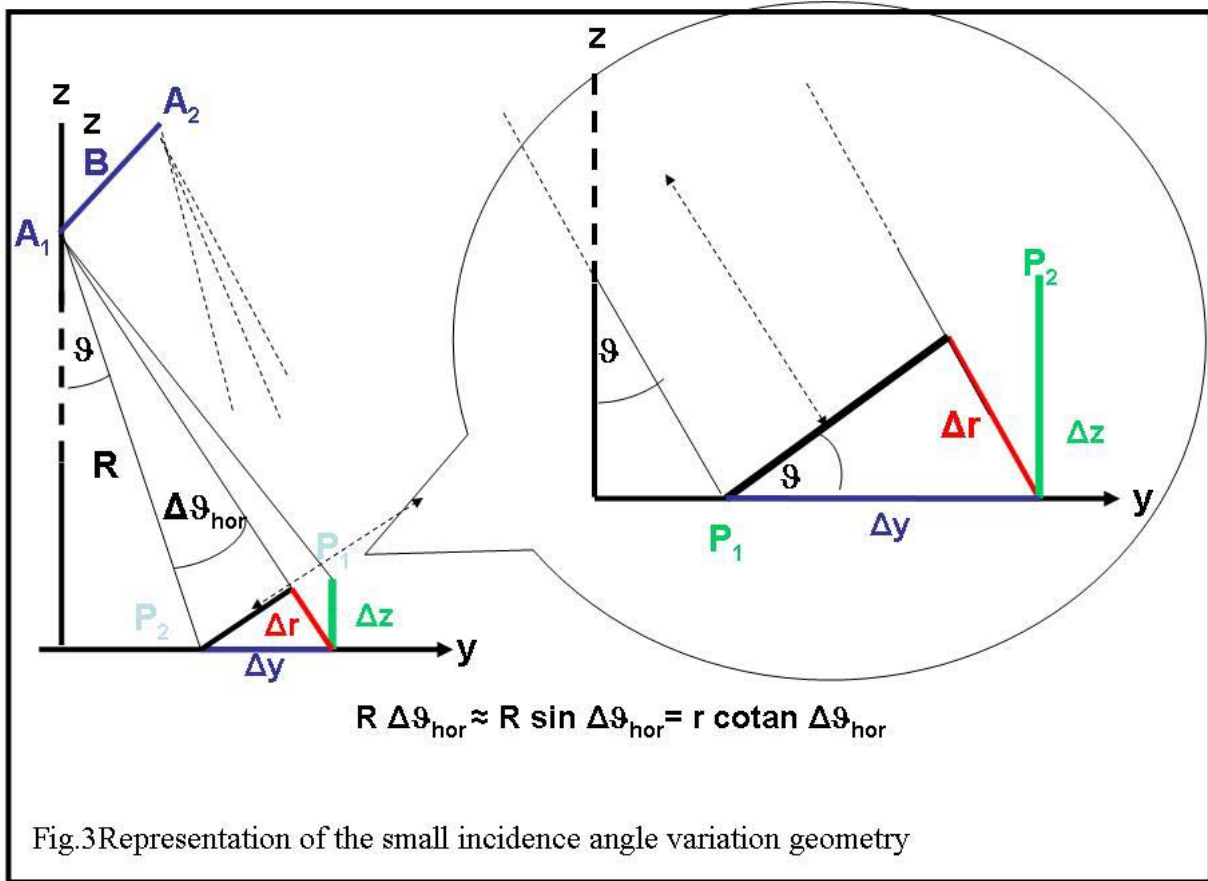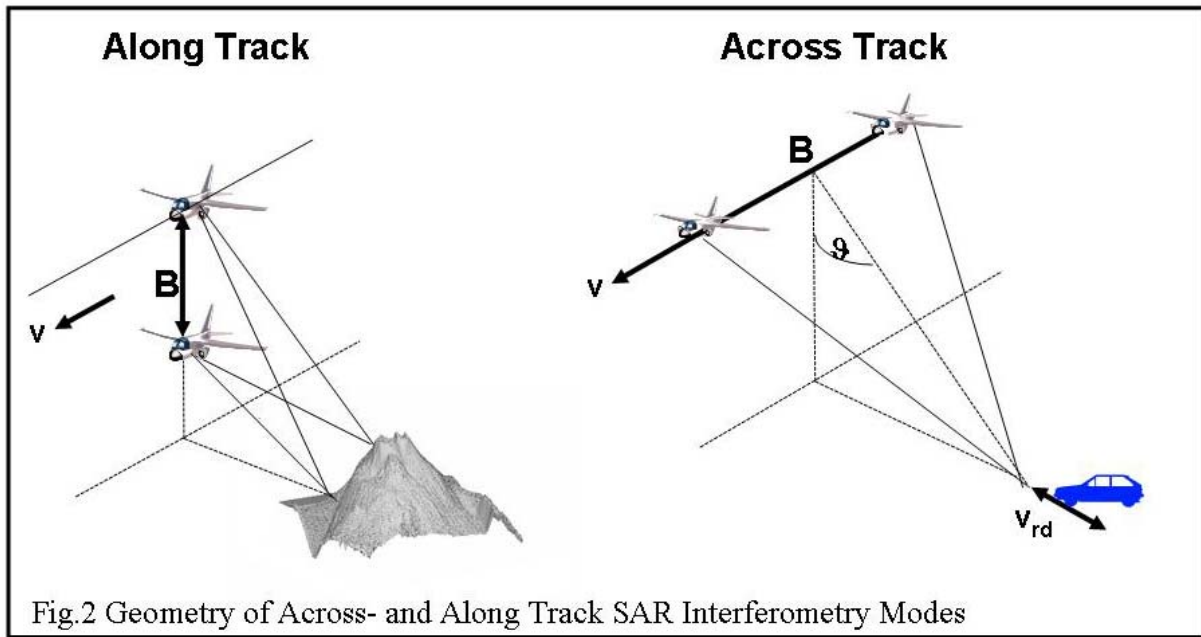
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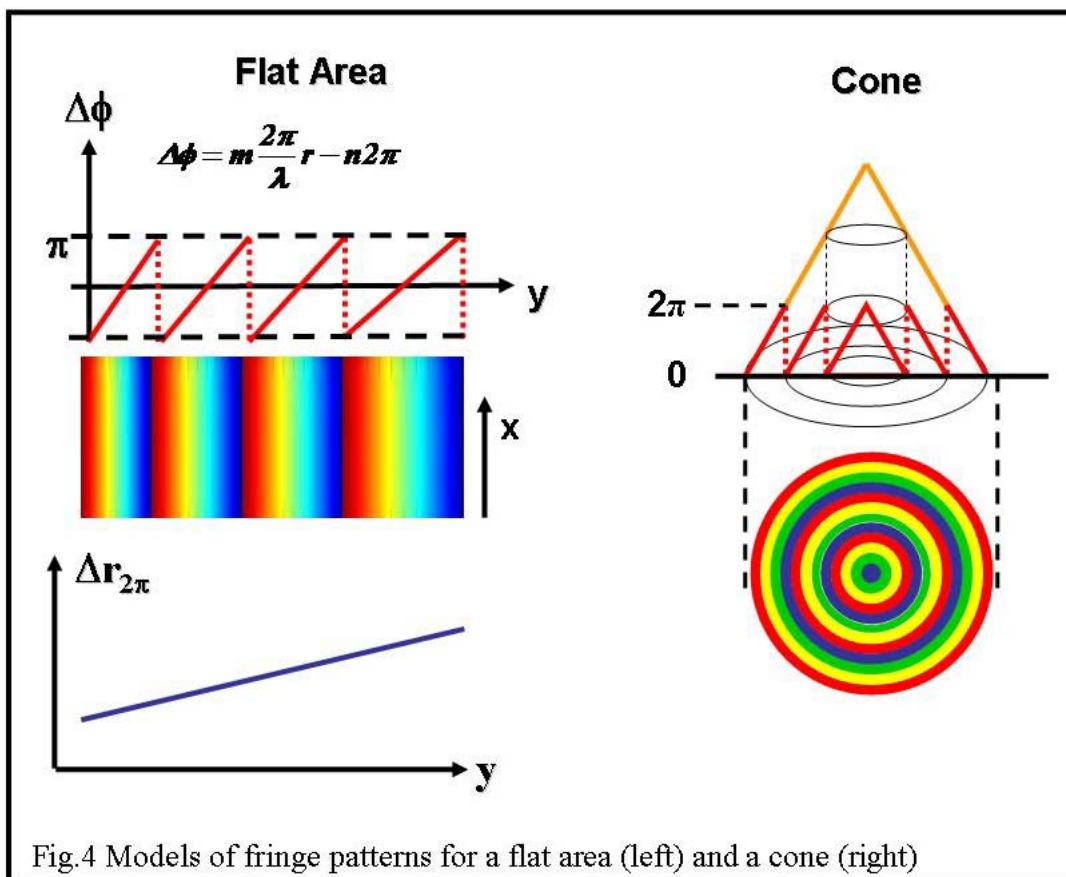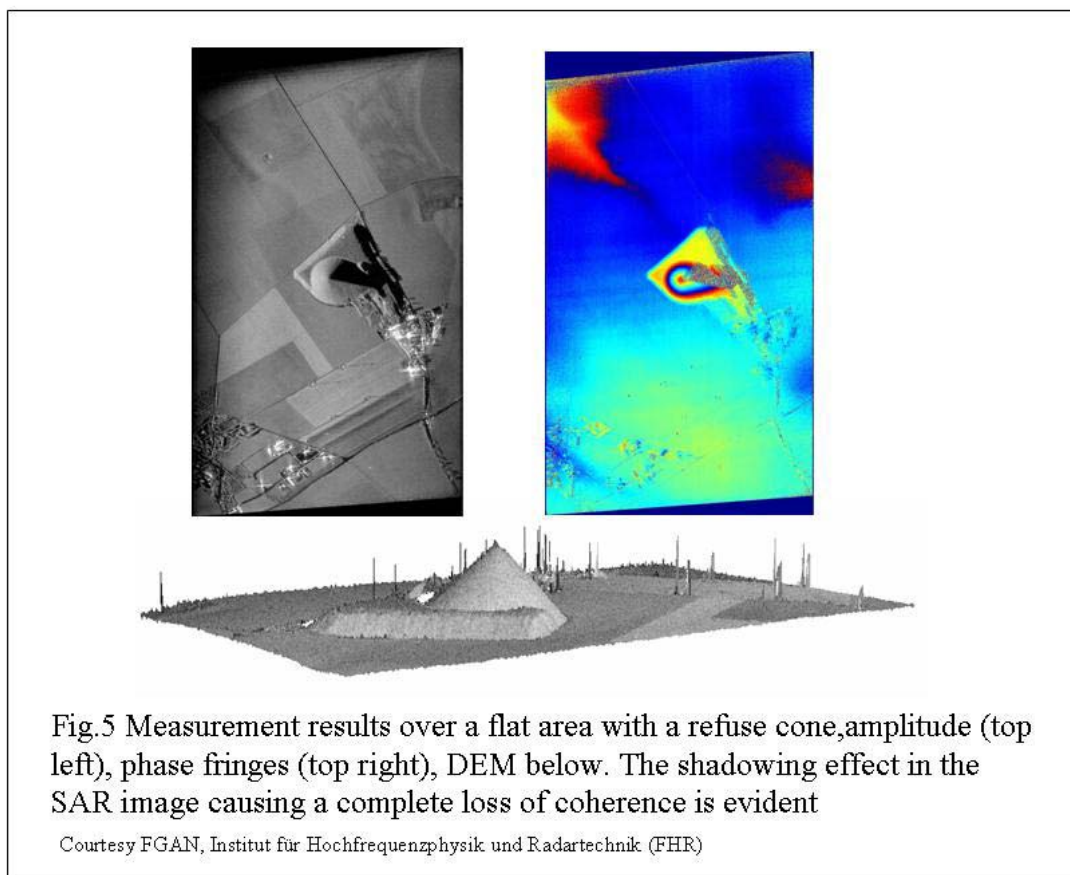
10. Figures

Fig. 1 SAR Interferometry Geometry

Fig.2 Geometry of Across- and Along Track SAR Interferometry Modes

$$R \, \Delta\vartheta_{hor} \approx R \sin \Delta\vartheta_{hor} = r \cot \Delta\vartheta_{hor}$$

Fig.3 Representation of the small incidence angle variation geometry

Flat Area

$$\Delta\phi = m\frac{2\pi}{\lambda}r - n2\pi$$
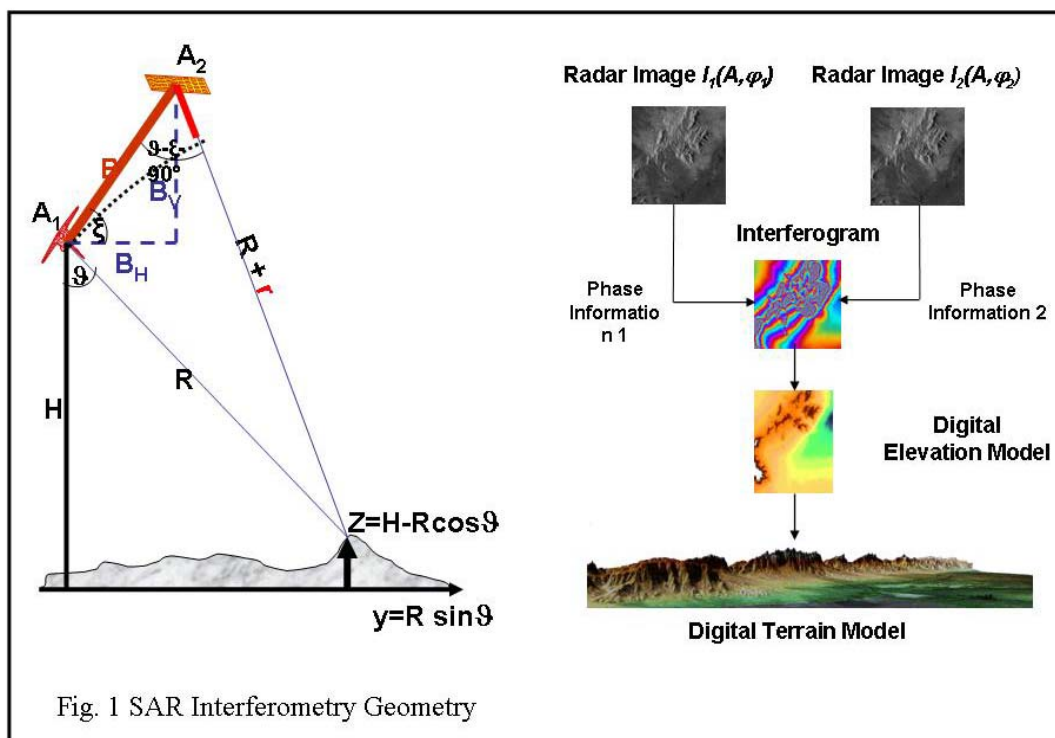
Fig.4 Models of fringe patterns for a flat area (left) and a cone (right)


Fig.5 Measurement results over a flat area with a refuse cone, amplitude (top left), phase fringes (top right), DEM below. The shadowing effect in the SAR image causing a complete loss of coherence is evident

Courtesy FGAN, Institut für Hochfrequenzphysik und Radartechnik (FHR)
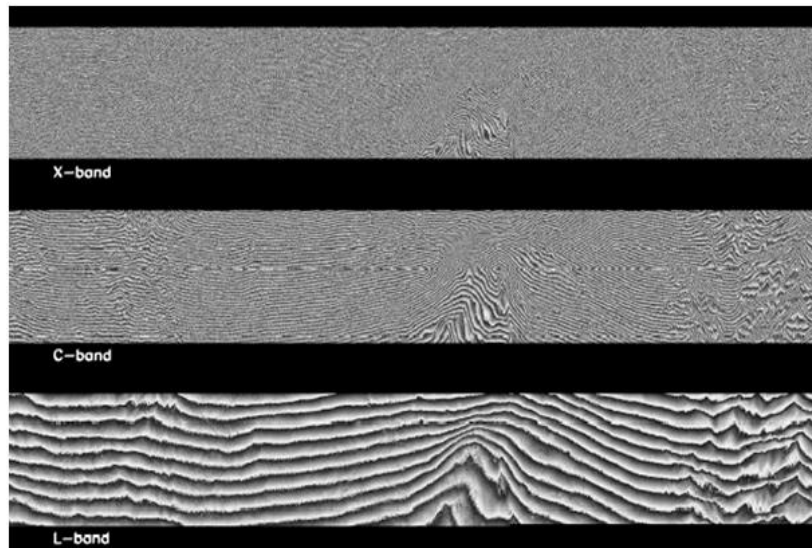
Fig 6 Set of phase fringes without flat earth removal, taken simultaneously in X-, C-, and L-Band with a two pass Interferometry experiment with SIR-C over the Mount Etna Test area. Time difference between the two passes was one day. The Interferogram after flat earth correction is depicted in Fig 17. The area in the left part was relatively flat, the largest slopes of the volcano is nearly in the middle. The image size is 66.8 km in azimuth and 9 km in ground range (16000 radar pulses and 1024 range bins in X-band and 2048 range bins in L- and C-band).
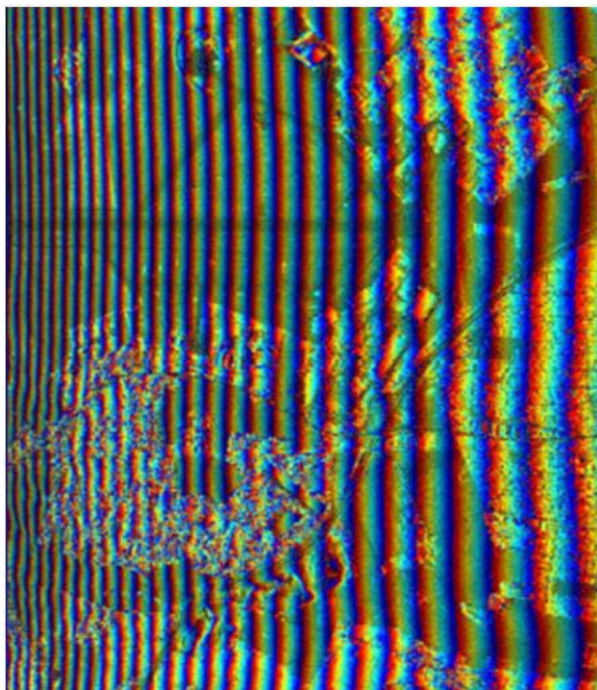
Near Range	Far Range
$\vartheta = 25°$	$\vartheta = 55°$
R = 3530 m	R =5579 m
B = 28.1 m	11,5 m
2π ~ 10,3m Height	2π ~ 49,6m Height
2π ~ 26,9 m Ground Range	2π ~ 105,7 m Ground Range

Fig.7 Example of an airborne interferogram, taken with E-SAR in L-band repeat-pass mode in the relatively flat area around Oberpfaffenhofen/Germany therefore, mainly flat-earth fringes are visible only. Their frequency changes strongly from near-range to far range
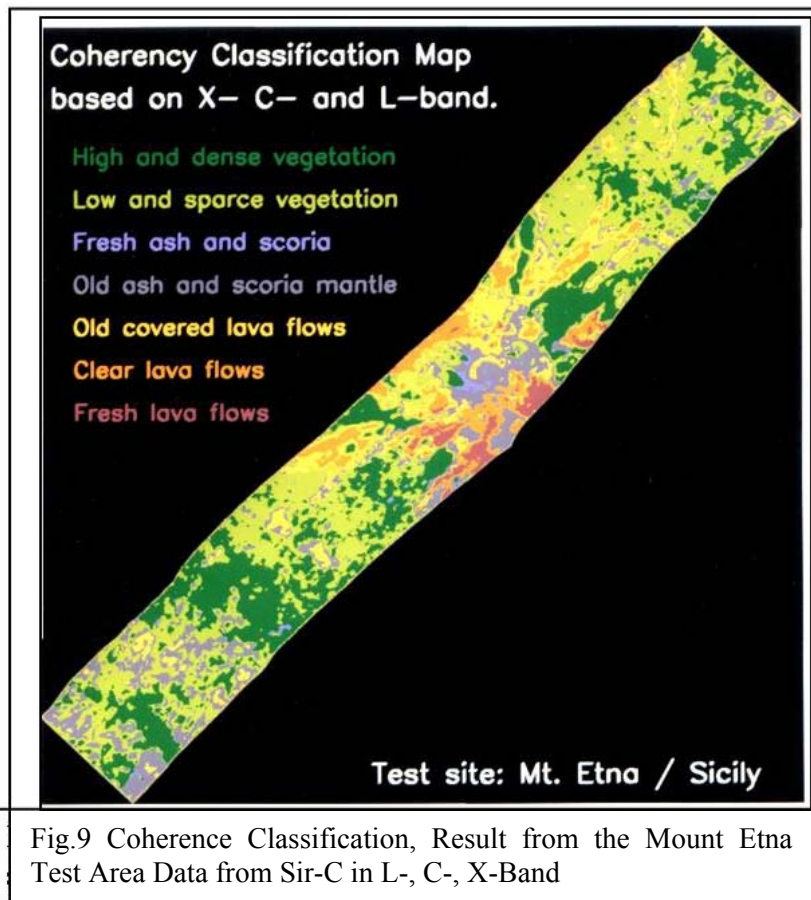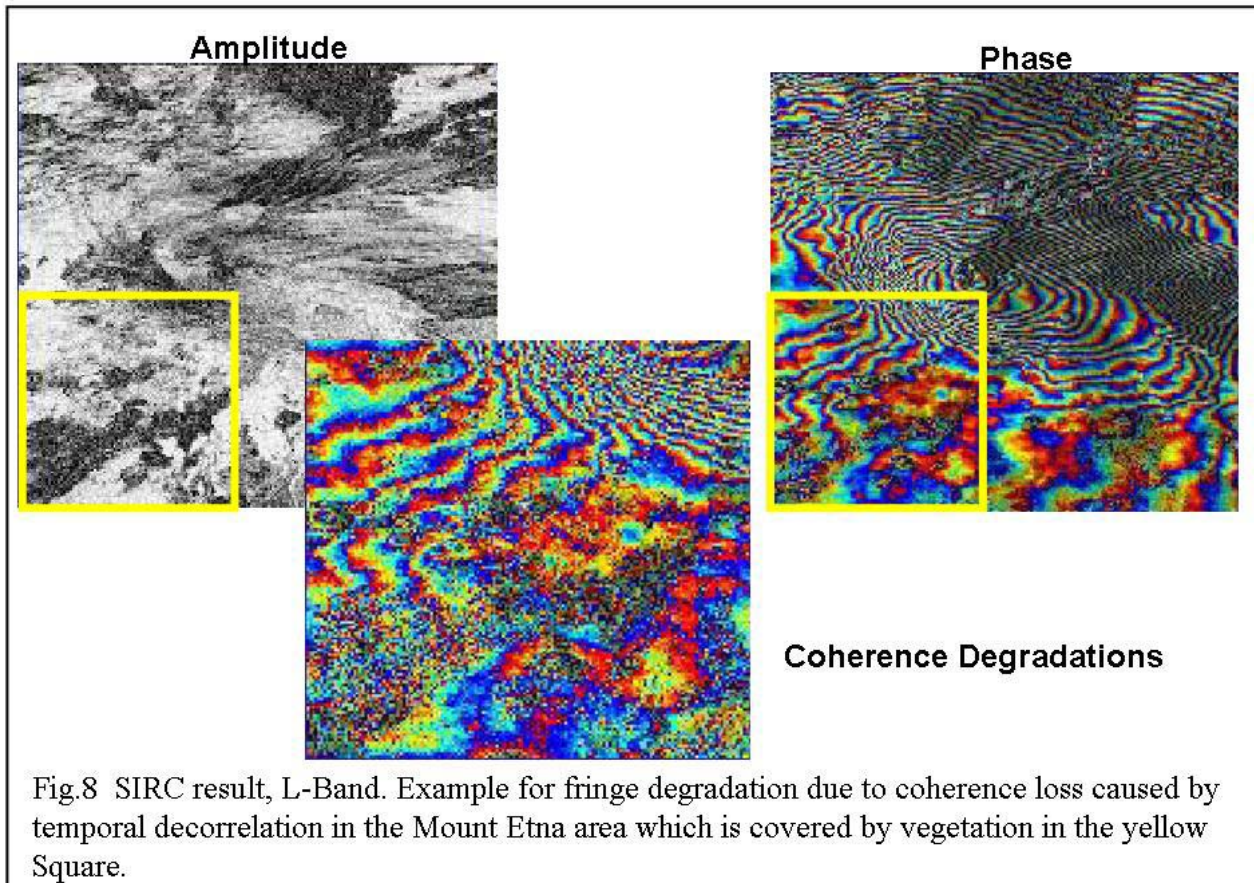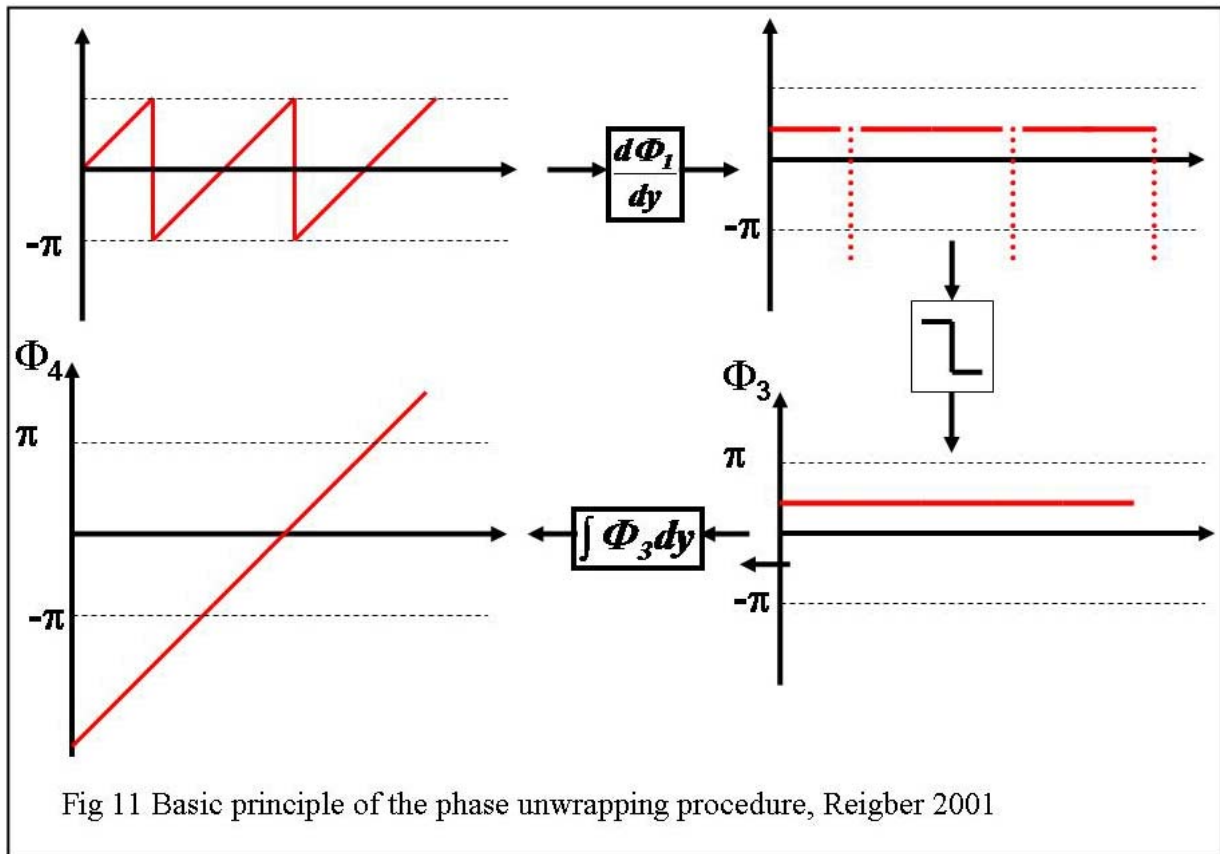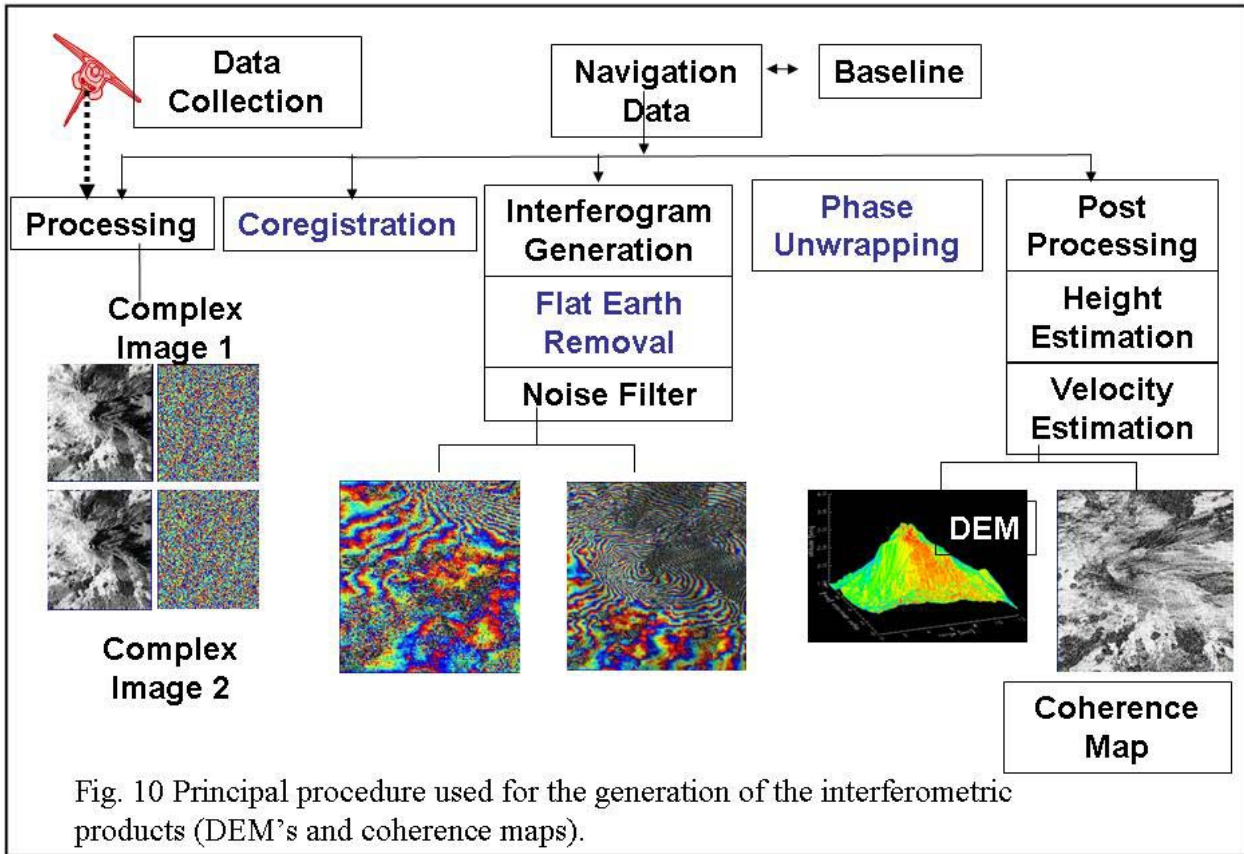
Fig.8 SIRC result, L-Band. Example for fringe degradation due to coherence loss caused by temporal decorrelation in the Mount Etna area which is covered by vegetation in the yellow Square.
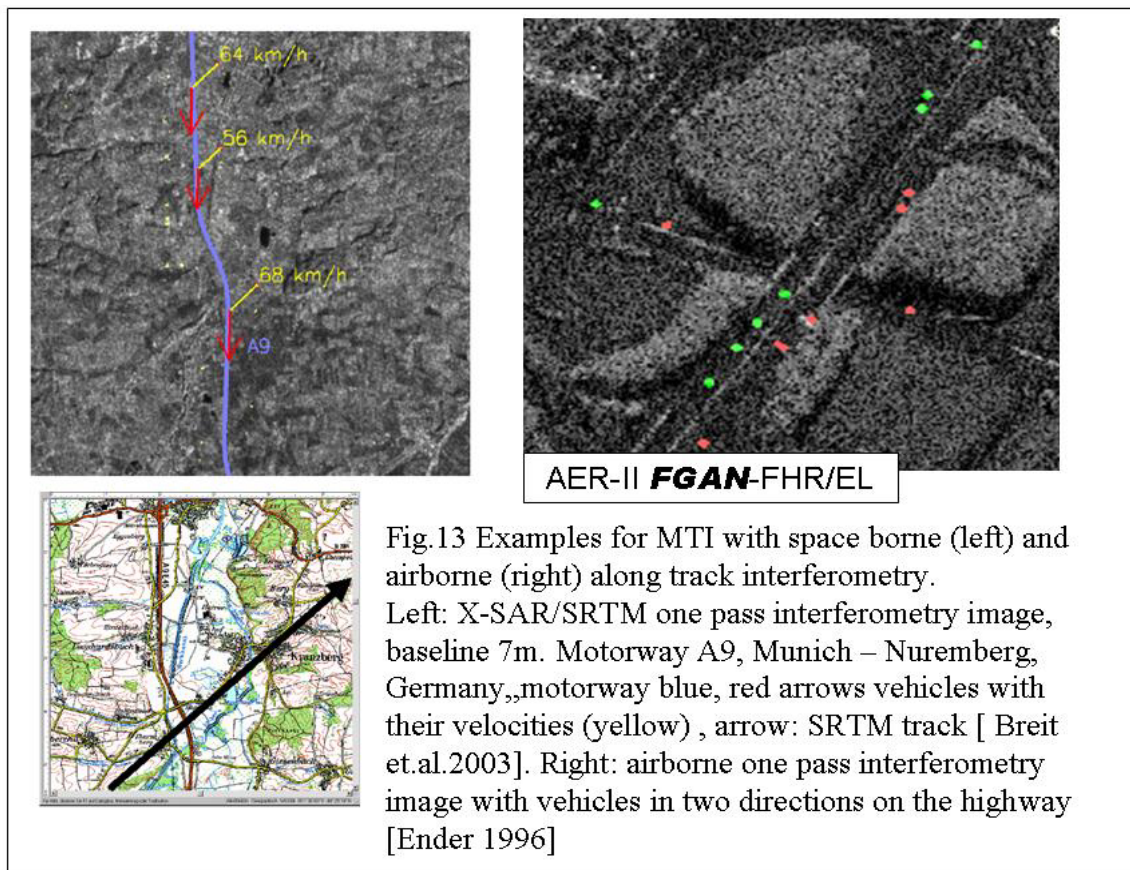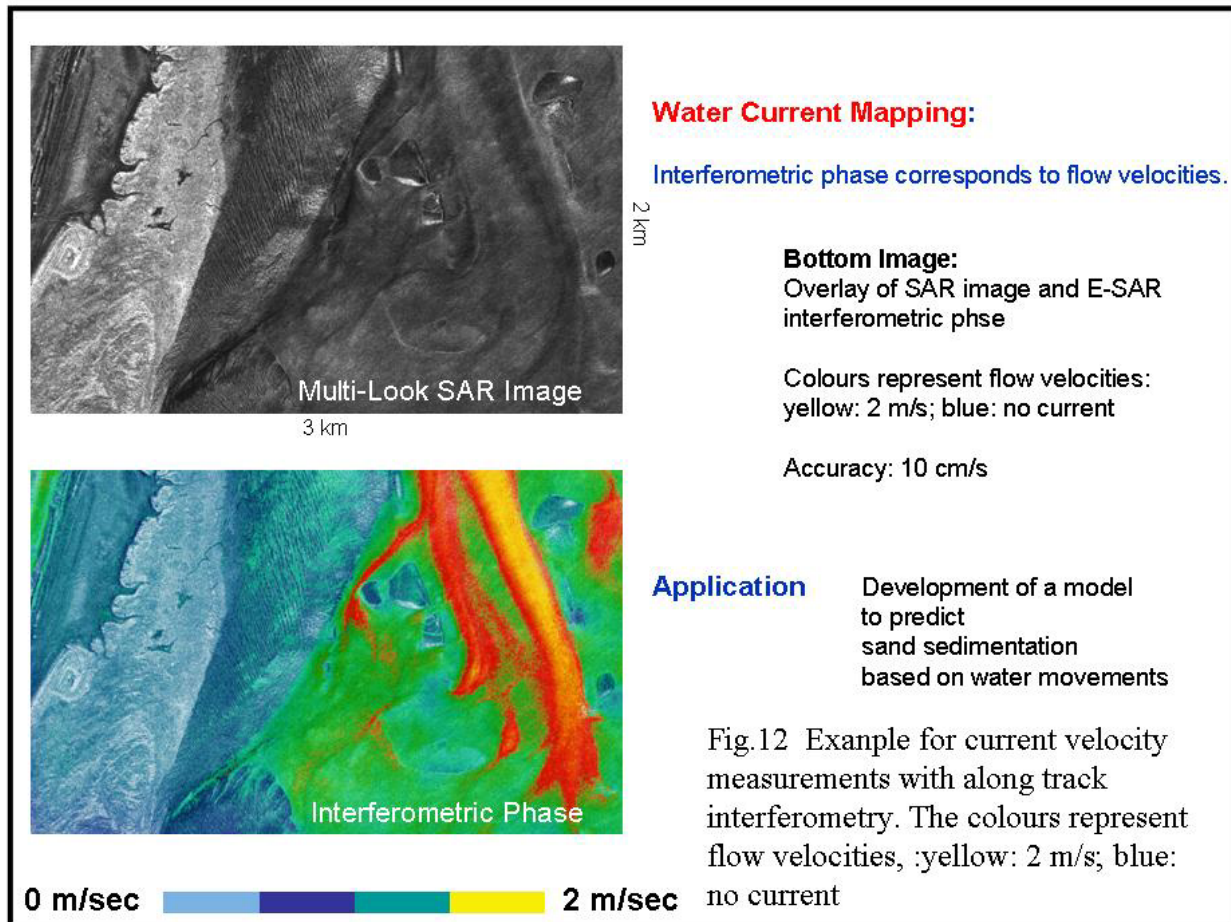

Fig.9 Coherence Classification, Result from the Mount Etna Test Area Data from Sir-C in L-, C-, X-Band

Fig. 10 Principal procedure used for the generation of the interferometric products (DEM's and coherence maps).

Fig 11 Basic principle of the phase unwrapping procedure, Reigber 2001

Water Current Mapping:

Interferometric phase corresponds to flow velocities.

Bottom Image:
Overlay of SAR image and E-SAR interferometric phse

Colours represent flow velocities:
yellow: 2 m/s; blue: no current

Accuracy: 10 cm/s

Application Development of a model to predict sand sedimentation based on water movements

Fig.12 Example for current velocity measurements with along track interferometry. The colours represent flow velocities, :yellow: 2 m/s; blue: no current

AER-II *FGAN*-FHR/EL

Fig.13 Examples for MTI with space borne (left) and airborne (right) along track interferometry.
Left: X-SAR/SRTM one pass interferometry image, baseline 7m. Motorway A9, Munich – Nuremberg, Germany„motorway blue, red arrows vehicles with their velocities (yellow) , arrow: SRTM track [Breit et.al.2003]. Right: airborne one pass interferometry image with vehicles in two directions on the highway [Ender 1996]

Differential Interferometry

Displacement Sensitivity

$$\Delta\rho = \frac{\lambda}{m2\pi}\Delta\Phi_2$$

$$\Delta\Phi_2 = m\frac{2\pi}{\lambda}\Delta\rho$$

$$\Delta z = \Delta\rho\cos\vartheta$$

Height Variation Sensitivity

$$\Delta\Phi_2 = m\frac{2\pi}{\lambda}\frac{\Delta h}{\cos\vartheta}$$

Fig.14 Differential Interferometry Geometry

Izmit/Turkey, 17 August, 1999

Data ERS-1/2 © ESA,
Acqusition dates: 12.08. and 16.09.1999

Izmit

Fig.15 Differential interferometry result., displacement pattern caused by an earthquake [Bamler, 2000]

Fig 16 Strucure of DLR's overall PS System, each box illustrates a sub-processor system will be applicated to each single input image separately. (Adam
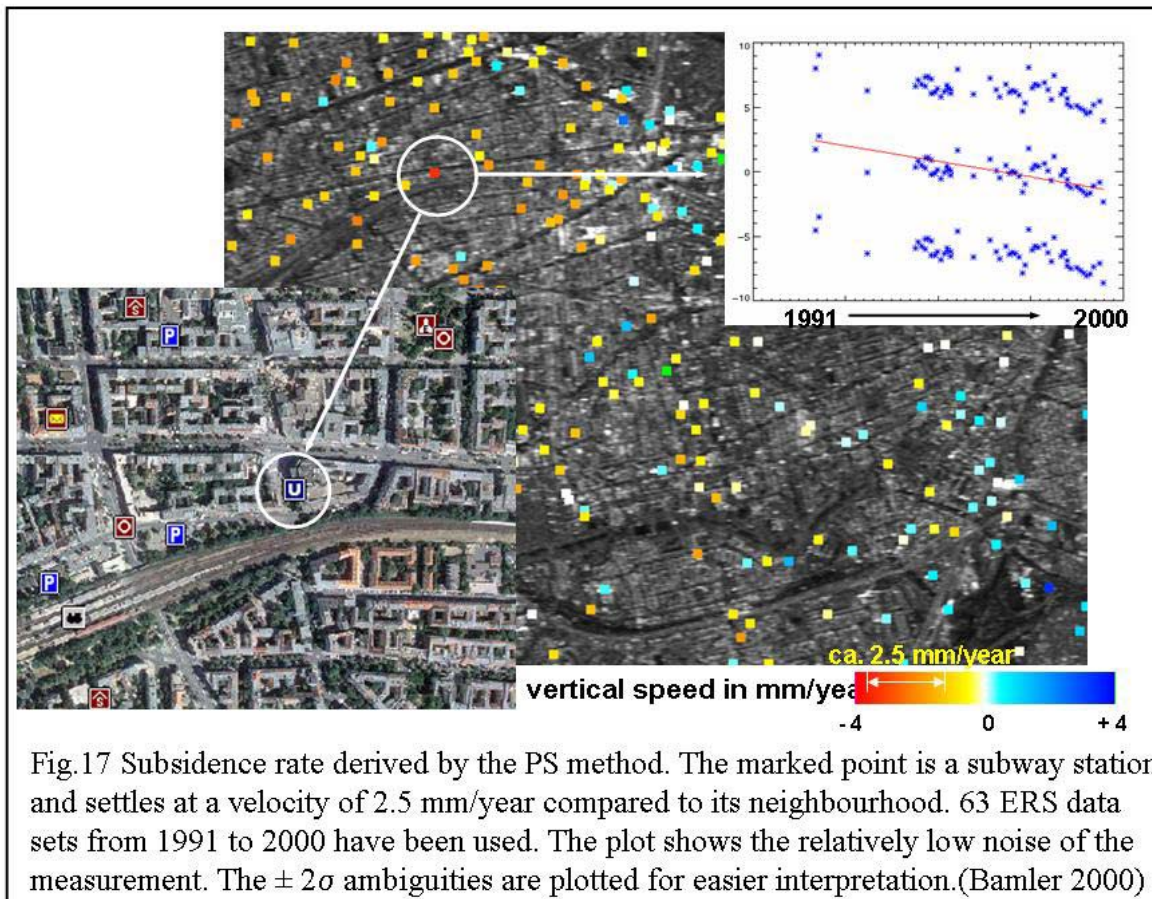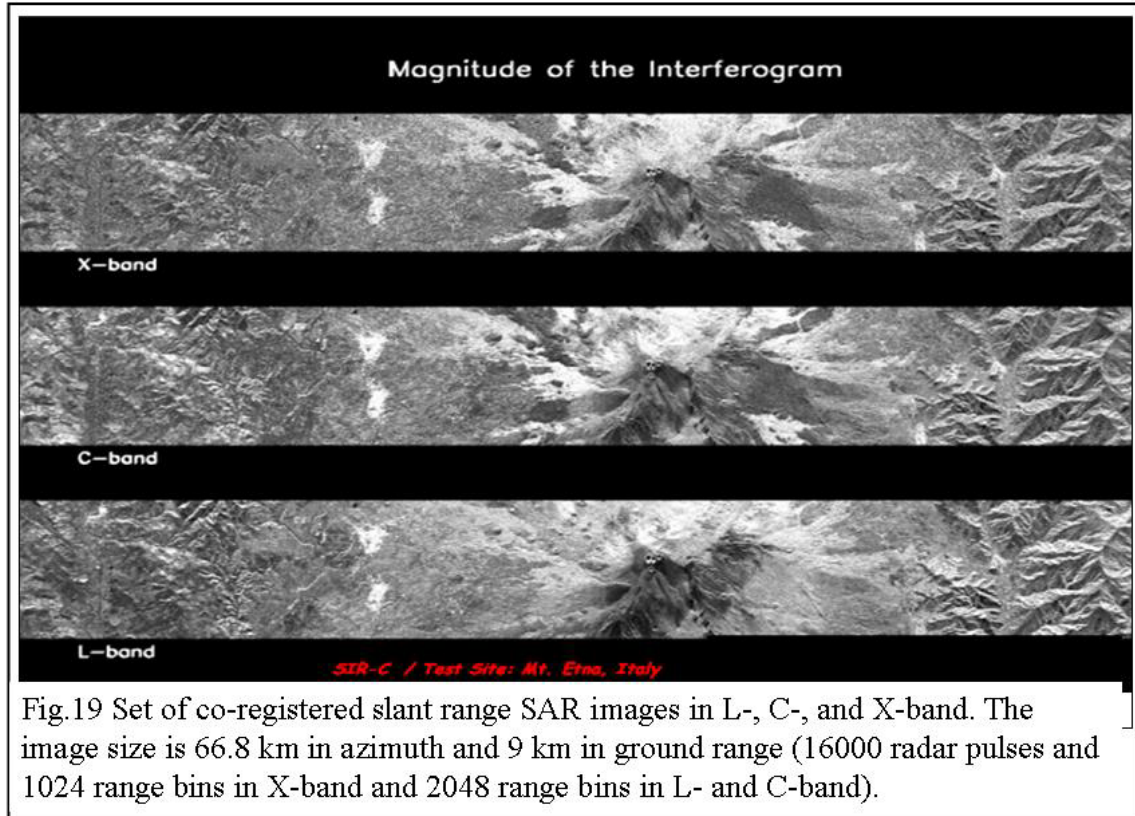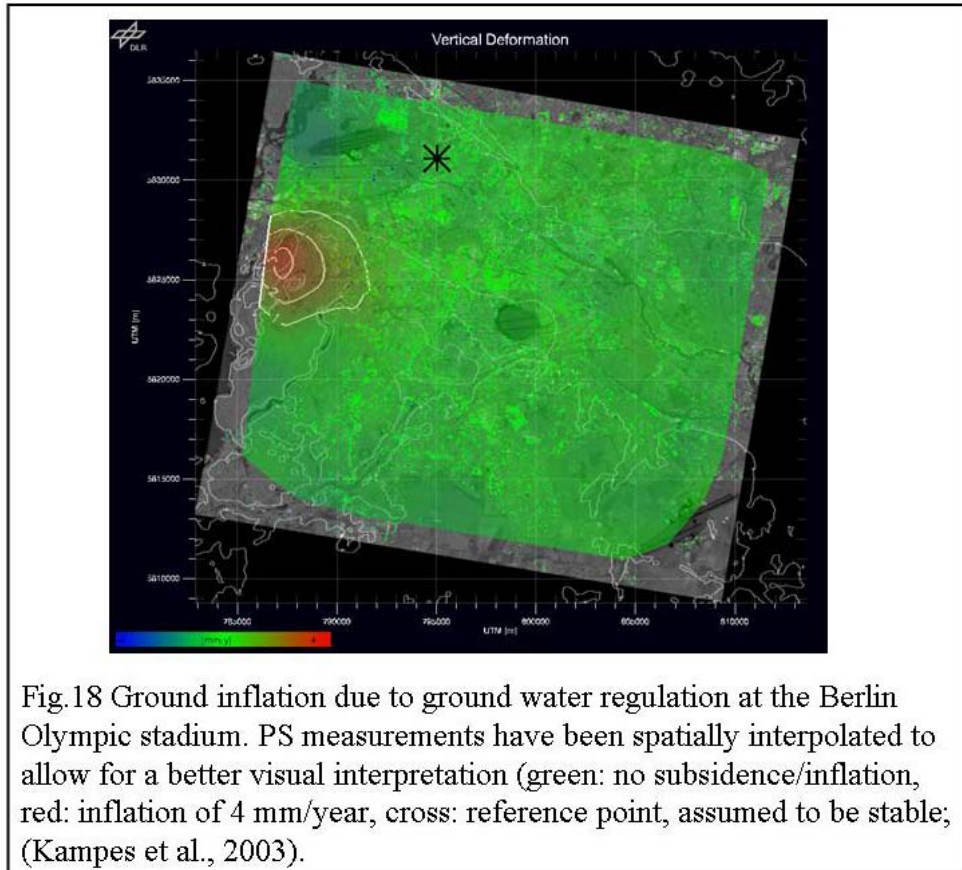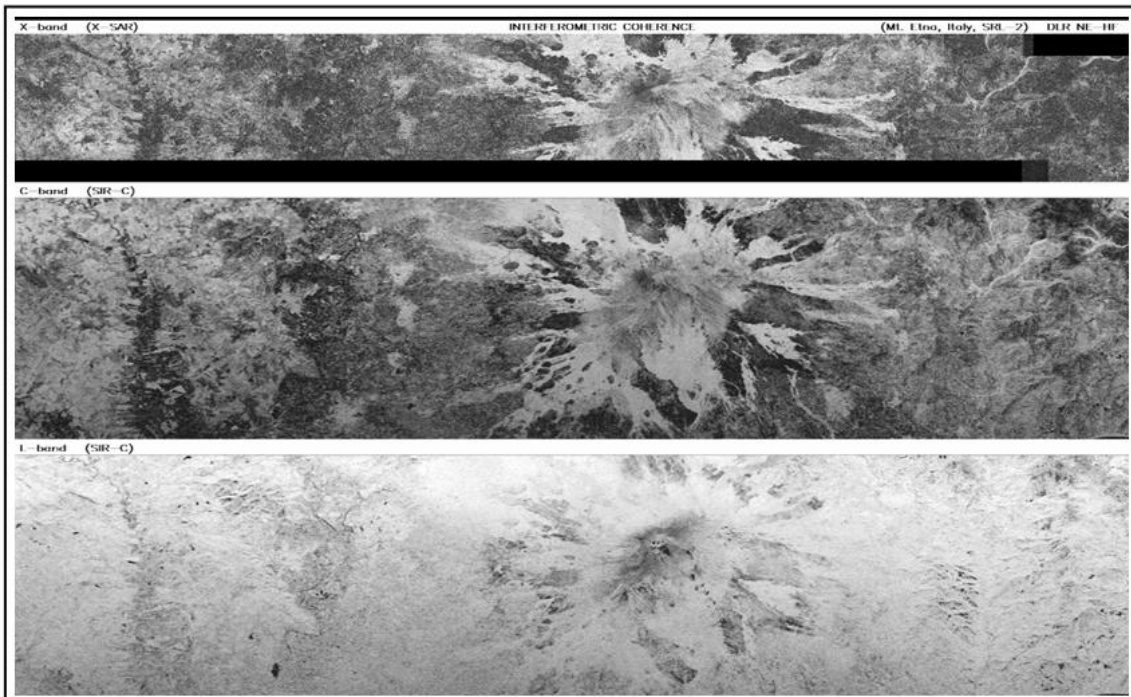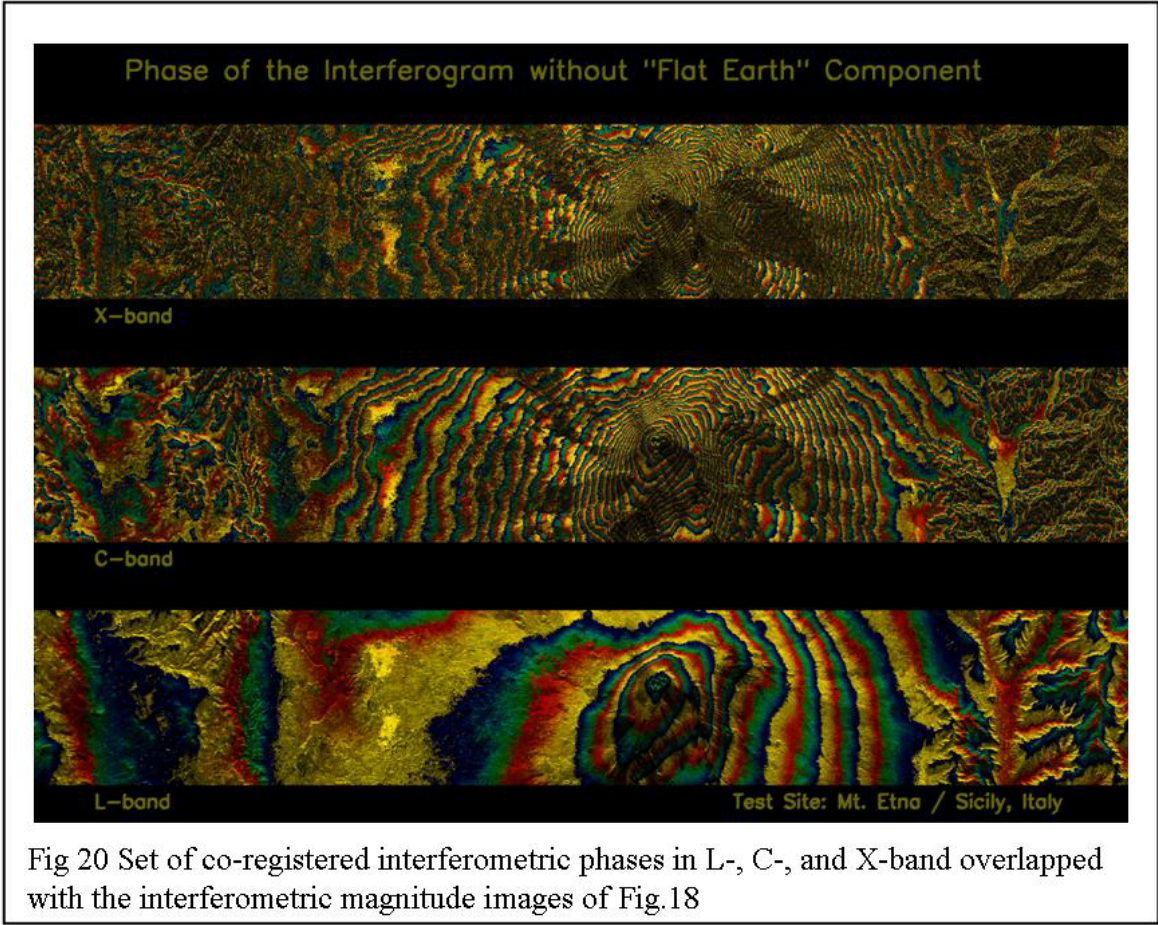

Fig.17 Subsidence rate derived by the PS method. The marked point is a subway station and settles at a velocity of 2.5 mm/year compared to its neighbourhood. 63 ERS data sets from 1991 to 2000 have been used. The plot shows the relatively low noise of the measurement. The $\pm 2\sigma$ ambiguities are plotted for easier interpretation.(Bamler 2000)
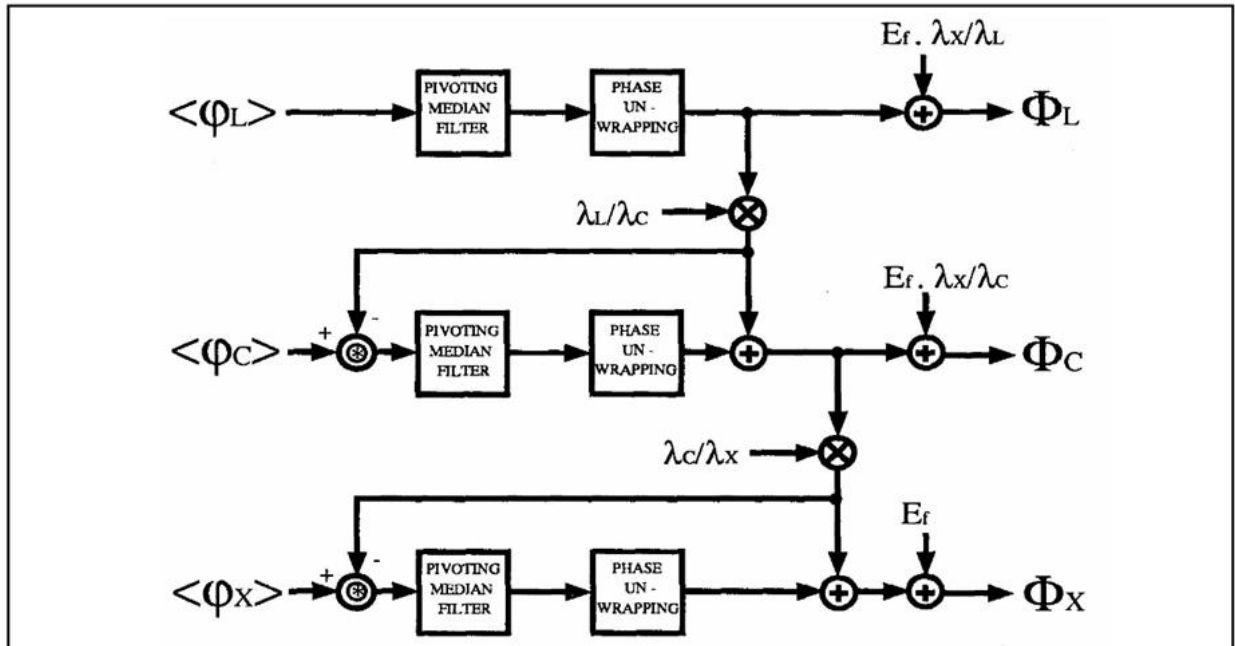
Fig.18 Ground inflation due to ground water regulation at the Berlin Olympic stadium. PS measurements have been spatially interpolated to allow for a better visual interpretation (green: no subsidence/inflation, red: inflation of 4 mm/year, cross: reference point, assumed to be stable; (Kampes et al., 2003).

Fig.19 Set of co-registered slant range SAR images in L-, C-, and X-band. The image size is 66.8 km in azimuth and 9 km in ground range (16000 radar pulses and 1024 range bins in X-band and 2048 range bins in L- and C-band).

Fig 20 Set of co-registered interferometric phases in L-, C-, and X-band overlapped with the interferometric magnitude images of Fig.18

Fig21 Set of co-registered slant range coherence images in L-, C-, and X-band. White indicates highcoherence and black low coherence. It is evident that l-Band has ahigh coherence with the exception of certain small areas

FIG 22 Block diagram of the multi-frequency phase unwrapping procedure. ($<\varphi L>$, ($<\varphi C>$ and $<\varphi X>$ are wrapped phases of the L-, C-, and X-band, respectively, without flat-earth component. λL λC λX are wavelengths. *Ef* flat-earth function scaled for X-band used in L- and C-band with the respective wavelength scaling. * denotes phase difference calculated in the complex domain, $\Phi L \Phi_C \Phi_X$, are unwrapped original phases in the respective Bands
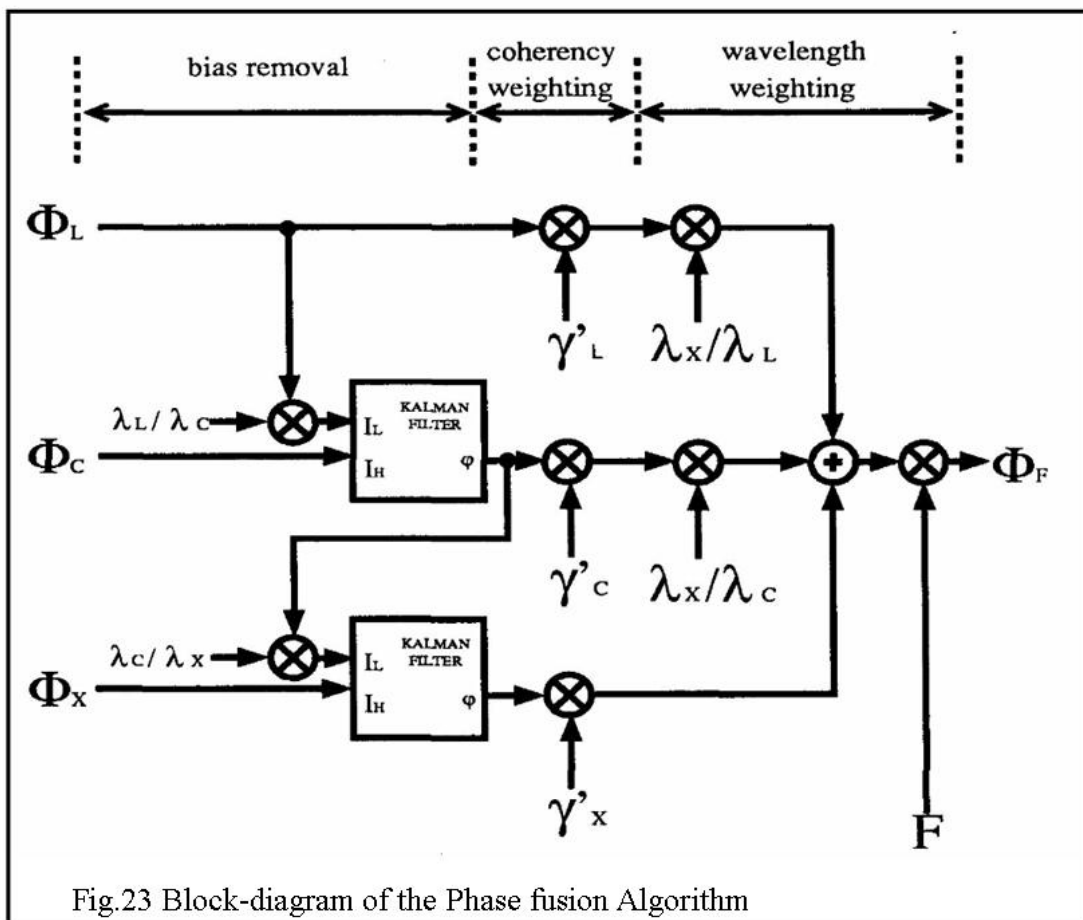
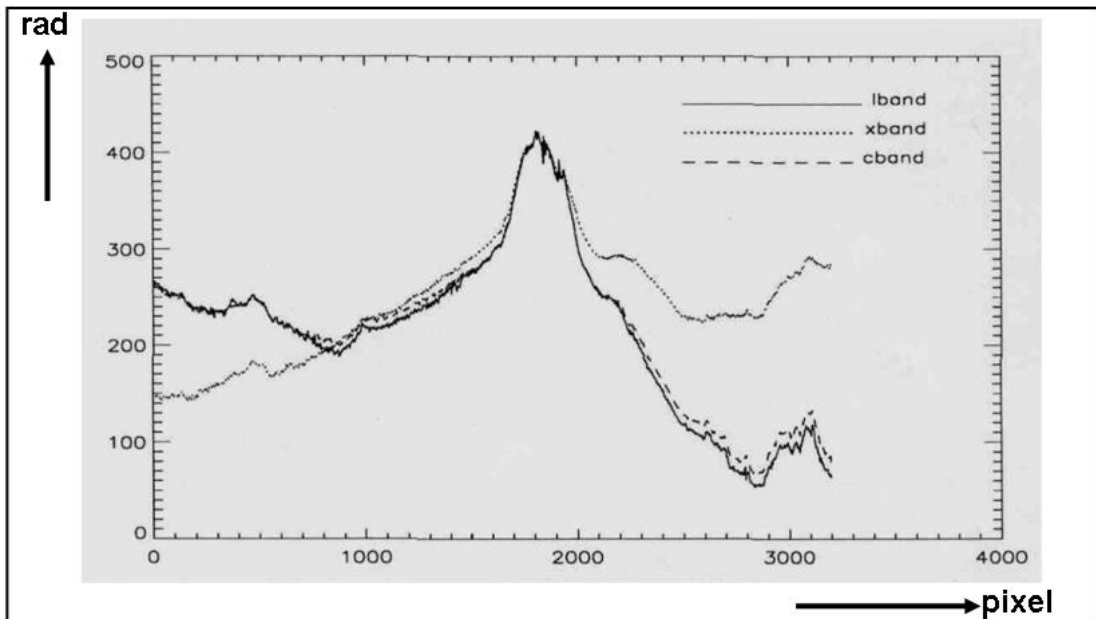Fig.23 Block-diagram of the Phase fusion Algorithm

Fig. 24 Unwrapped phase profiles(azimuth direction) filtering without unwrapped phase profiles of L-, C-, and X-band in azimuth direction applying multi-frequency phase unwrapping without filter before phase unwrapping. The L- and C-band phases were scaled with respect to the X-band wavelength for comparison purposes. The x-axis gives the azimuth dimension in pixels. Each pixel has a size of 20.9 m (3200 pixels correspond 66.864 km). The y-axis unit is radians.
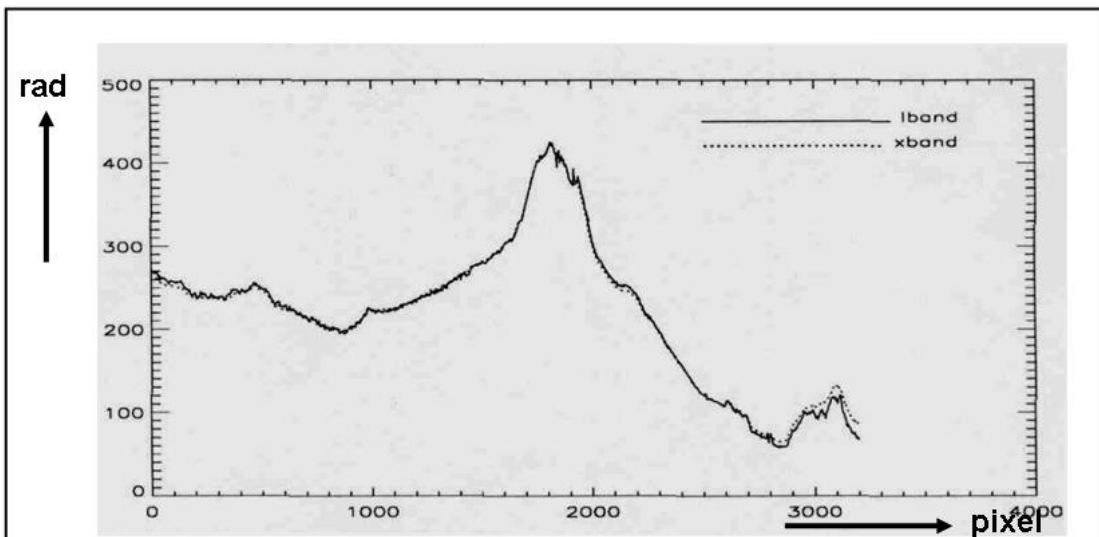

Fig.25 Unwrapped phase of L-, and X-band profiles in azimuth direction applying multi-frequency phase unwrapping with the cascade of two filters with a window of 5x5 before phase unwrapping. The L- phase was scaled with respect to the X-band wavelength for comparison purposes. The x-axis gives the azimuth dimension in pixels. Each pixel has a size of 20.9 m (3200 pixels correspond to 66.864 km). The y-axis unit is radians.
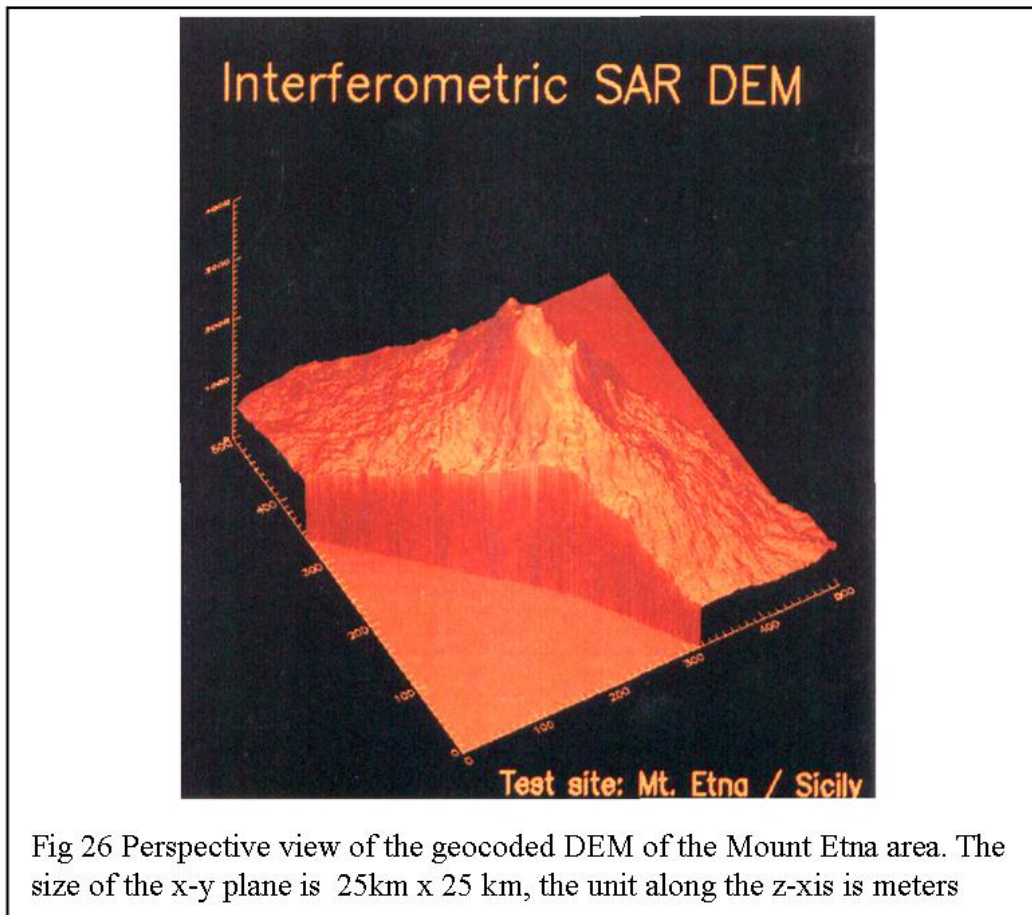
Fig 26 Perspective view of the geocoded DEM of the Mount Etna area. The size of the x-y plane is 25km x 25 km, the unit along the z-xis is meters